

University of Nevada, Reno

A Model for Simulating Soil-Zone Processes at the Regional Scale

A thesis submitted in partial fulfillment of the
requirements for the degree of Master of Science in
Hydrogeology

by

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We recommend that the thesis
prepared under our supervision by

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A Model For Simulating Soil-Zone Processes At The Regional Scale

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ABSTRACT:

The soil zone has been described as the upper most region of the vadose zone where plant and soil processes enhance storage and permeability, providing a fast pathway for water and solutes to streams. The soil zone connects the land surface to the deeper unsaturated and saturated zones. Modeling of soil-zone processes has been used to gain understanding of watershed hydrologic processes. Currently MODFLOW does not simulate dynamic near-surface hydrologic processes such as, infiltration, hortonian runoff, dunnian runoff, and return flow. The Soil-Zone Flow Package (SZF) for MODFLOW is being developed to address these near-surface components for simulating watershed processes in the context of basin-scale groundwater-flow modeling. especially those processes that partition rainfall into evapotranspiration, runoff, and deep percolation.

In a series of test simulations, Richards' equation (RE) was compared with MODFLOW-SZF. Across a range of hydraulic conductivities and applied precipitation rates, MODFLOW-SZF achieved a good infiltration and interflow solution (0.014-0.052 RMSE) with only 6 grid cells and 0.11 seconds computation time compared to the 6250 grid cells and 40 seconds of computation time required for a stable solution to RE. MODFLOW-SZF solutions had negligible errors due to grid effects (<0.01 RMSE) however, RE solutions were sensitive to grid resolution (up to 0.12 RMSE). Errors associated with using the groundwater flow equation to represent soil-zone flow instead of RE were much lower (<0.052 RMSE) than the errors caused by grid effects in RE (0.12 RMSE, 0.5 m cell). This research suggests that the SZF Package will be an effective tool for efficiently representing soil-zone processes at the basin scale.

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CHAPTER 1: INTRODUCTION

Numerical models that use process based physics are needed to better understand the effects of water resources development, climate change, and hydrologic controls on ecosystems (Harter and Hopmans, 2005). Departing from the traditional view of groundwater modeling as only simulating the baseflow component of a hydrograph, The Soil-Zone Flow (SZF) Package for MODFLOW can simulate the entire streamflow hydrograph, including runoff, interflow, and baseflow using spatially distributed parameters and physics-based equations.

This thesis presents a new enhancement to MODFLOW (SZF Package) for simulating near-surface hydrologic processes coupled to the three-dimensional groundwater-flow equation that is solved by MODFLOW. Additionally, this thesis describes the conceptualization of the SZF Package and presents comparisons to another numerical model that has been used to simulate watershed flow processes.

PROBLEM STATEMENT

The soil zone partitions precipitation into runoff, interflow, ET, and deep percolation. Understanding the role of the soil zone in groundwater recharge can help improve hydrologic models. The SZF Package for MODFLOW will simulate soil-zone processes using physics-based equations by partitioning precipitation into runoff, infiltration, evapotranspiration, subsurface-storm flow (interflow), and deep percolation.

Although the SZF Package simulates all of these near-surface hydrologic processes, this thesis focuses on evaluating the interflow component of watershed drainage.

PREVIOUS WORK

IMPORTANCE OF SOIL-ZONE FLOW

The permeability of the soil zone typically is enhanced relative to underlying sediments and rocks due to biologic activity and root zone processes. Chemical leaching from the soil zone may also reduce the permeability of sediment and rock beneath the soil zone, further increasing the contrast in permeability. The vadose zone, which often includes the soil zone, regulates flow to and from groundwater (Harter and Hopmans, 2005). Soils serve several important functions. Soils act like sponges soaking up rainwater and limiting runoff, which in turn affect deep percolation, groundwater recharge and flood-control potential in urban and natural areas. Soils store and release water, air, and nutrients for plants and animals to use. Soils sequester carbon, filter water, as well as buffer, degrade, immobilize, detoxify, and trap pollutants (Scheyer & Hipple, 2005). More than 95% of water in streamflow has passed over or through soils before reaching the channel network (Kirkby, 1988).

WATERSHED HYDROLOGY

Watershed hydrology is the study of hydrologic processes that occur within watersheds such as infiltration, runoff, interflow, and subsurface flow. Knowledge of relevant processes is integral to understanding watershed hydrology (Beven, 1989). Therefore, simulation of soil-zone processes such as infiltration, runoff and interflow will improve the accuracy of basin-scale hydrologic models. The SZF Package will simulate these processes.

Developing an understanding of watershed processes is important in the context of basin-scale modeling. The relative importance of each hydrologic process in a particular region is affected by watershed characteristics, including the type of climate, geology, physiography, soils, vegetation and land use (Dunne, 1978). For example, regions with semi-arid to arid climates, runoff processes are less common, but in wetter climates these processes dominate watershed drainage. In order to simulate watershed-hydrologic processes, it is important to delineate how watershed characteristics vary with time, and space. The SZF Package is a physically-based distributed parameter model that considers areally distributed soil-zone properties and climate forcing to consider spatial variability in climate, geology, physiography, soils, vegetation and land use.

A conceptual cross section of a watershed (fig. 1) is divided into three hydrogeologic layers: a soil-zone layer, an unsaturated zone layer, and a groundwater layer. This thesis will focus on processes occurring in the soil zone. Precipitation is aerially distributed across the watershed and collects on the surface. Once precipitation (rain) contacts the soil surface, the processes driving infiltration and runoff generation

begin. Precipitation is then partitioned into runoff and infiltration. Evapotranspiration (ET) of soil moisture is removed from the soil zone by vegetation and bare-soil evaporation. Infiltration and runoff are functions of potential ET, antecedent-soil conditions, soil type, soil composition, degree of vegetation, and macropore distribution. The relationship between runoff and infiltration directly impacts how precipitation is partitioned between the surface and subsurface, determining how water is transmitted between the surface and subsurface layers such as the soil zone, unsaturated zone, and groundwater.

Groundwater perches in the soil zone when deep percolation is less than infiltration at land surface. When groundwater perches in the soil zone, horizontal flow occurs downslope along the restrictive layer (interflow) and is routed to a stream channel or percolates downward to the regional aquifer where the restrictive layer is absent (Scheyer & Hipple, 2005; Dunne, 1978). Even though interflow is typically a temporary storm-related phenomena, it can represent a large component of watershed drainage in watersheds. The proportion of streamflow that is comprised of interflow has been estimated to be 2 to 34% of precipitation and 6 to 86% of total runoff, the average being 10 and 23%, respectively (Hewlett and Hibbert, 1967).

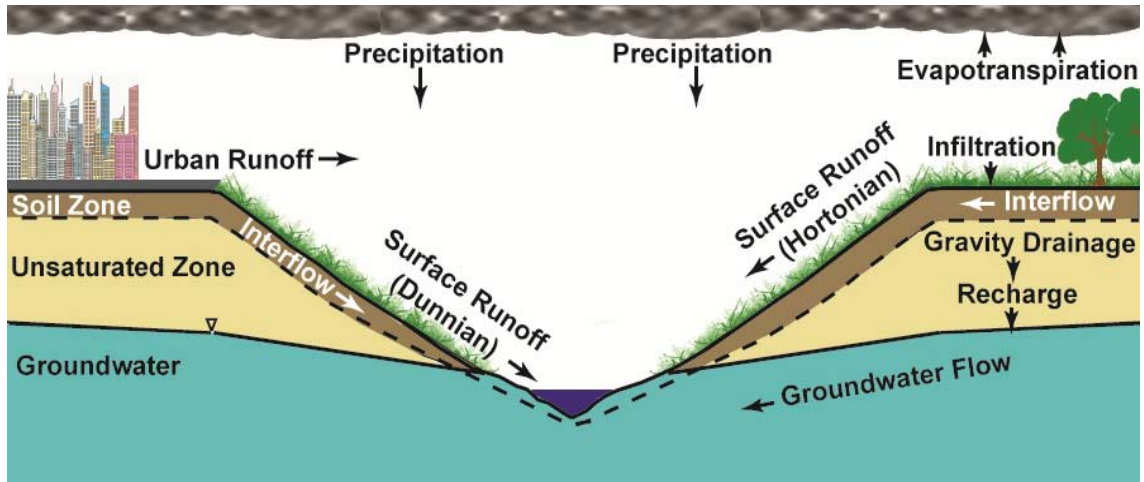


Figure 1: Conceptual watershed cross section

VARIABLY SATURATED FLOW

Richards (1931) derived an equation to describe unsaturated flow in three dimensions, using the Buckingham-Darcy formulation for unsaturated flow (Buckingham, 1907). Richards' equation calculates variably-saturated flow using a factor of proportionality relating the forces of gravity and capillary pressure gradients to the flow rate through porous media, called unsaturated hydraulic conductivity. Unsaturated hydraulic conductivity can be calculated as a function of either capillary pressure or water content. Celia et al. (1990) showed that the mixed form of Richards' equation with special treatment of the time derivative provided the best results for simulating variably-saturated flow. Many researchers have derived relationships between unsaturated hydraulic conductivity and water content or matric potential (e.g., Brooks and Corey, 1966; Mualem, 1976, Van Genuchten, 1980). The Brooks-Corey hydraulic conductivity function is used in the SZF Package to calculate the unsaturated hydraulic conductivity as

$$K(\theta) = K_s \left[\frac{\theta - \theta_r}{\theta_s - \theta_r} \right]^\epsilon,$$

where $K(\theta)$ is the unsaturated hydraulic conductivity as a function of water content, θ , K_s is the saturated hydraulic conductivity, θ_r is the residual water content, θ_s is the saturated water content, ϵ is the Brooks-Corey shape factor.

Currently there is no unified, universally accepted mathematical representation of regional flow or transport in the unsaturated zone (Harter and Hopmans, 2005). Richards' equation is the universally accepted mathematical model but it has less physical meaning at the regional scale because the constitutive relationships that are used in the model were established for samples analyzed in the laboratory (<1 m length scale) (Vogel and Ippisch, 2008; Harter and Hopmans, 2005). Although Richards' equation has mostly been verified at the field scale, it is the current standard in the absence of a regional scale mathematical model to represent variably-saturated flow in three dimensions. The validity of Richards' equation in the laboratory and small scale experiments does not guarantee that the same equation will describe the average behavior of a large scale system equally well (Mantoglou, 1992). Harter and Hopmans (2005) state that neither three dimensional numerical models based on field scale parameterization of Richards' equation nor upscaled stochastic equations are practical because of the complexities and computational requirements associated with numerical solutions of these equations. Further, Jensen and Mantoglou (1992) showed unsaturated-flow paths essentially become vertical after upscaling to large grid cells (100 m), providing precedence for using 1-D

unsaturated-zone flow equations for regional-scale simulations. Accordingly, this was the approach used in the development of the SZF Package for MODFLOW.

A major factor that contributes to the complexity of regional scale models is the large degree of spatial variability exhibited by soil formations. While we can discretize governing partial differential equations at a scale fine enough to account for local fluctuations in soil properties, adequate data is not available to represent intrinsic spatial variability (Montoglou, 1987). Length correlation scales in the soil zone are on the order of 5-10 cm (Jensen and Montoglou, 1992), infiltration solutions are stable in simulations with vertical discretizations on the order of 0.1-1cm (Downer and Ogden, 2003), and assumptions that define the spatial scale of Richards' equation break down at approximately 4 m or less (Vogel and Ippisch, 2008); therefore, creating a transient 3-D Richards' equation regional-scale model may require billions of model nodes, and the use of multiple computers to obtain results faster than real time (Harter and Hopmans, 2005).

Some models explicitly account for vadose-zone processes but the issue of field versus basin scale measurement and computational requirements for basin-scale models still persists. The SZF Package uses an analytical solution to Richards' equation (Smith and Parlange, 1978; Smith and Corradini, 1993) that does not require vertical discretization, reduces computation time, and increases numerical stability for regional scale models. Analytical models assume vertical homogeneity in the soil zone. However, this assumption is not typically limiting in regional-scale models due to a lack of data that describe soil layering that is laterally extensive to traverse an entire model cell.

INFILTRATION

Many models have been developed to simulate infiltration (Green and Ampt, 1911); Horton, 1933; Phillip, 1957; Smith and Parlange, 1978; and Smith and Cordini, 1993). Infiltration models typically consider both capillary and gravitational forces that affect the flow through unsaturated-porous media.

Infiltration is complex even under assumed conditions of homogeneity and uniform initial water content (Mein and Larson, 1973). Infiltration capacity is a function of the saturated hydraulic conductivity, capillarity of the unsaturated soil, initial water content, and the cumulative infiltration. Generally, the infiltration capacity is a function of rainfall characteristics such as intensity, duration, and drop size; and soil characteristics such as texture, structure, depth, and clay mineralogy (Dunne, 1978). Infiltration and runoff are interrelated by the time of ponding, which describes when the precipitation rate exceeds the infiltration capacity of the soil. Following the time of ponding, excess precipitation becomes runoff.

RUNOFF

Runoff can be modeled as sheet flow until it channelizes into riffles. When the rainfall intensity exceeds infiltration capacity, surface depressions are filled first. This is referred to as depression storage. Once depression storage has filled, excess precipitation then spills over into overland flow (Dunne, 1978). Two types of overland flow are Hortonian runoff and Dunnian runoff. Hortonian runoff was first described by Horton (1933) as rainfall or snowmelt rates that exceed the infiltration rate at the soil surface (Kirkby, 1988). Excess infiltration ponds on the surface, and eventually moves down slope to a stream channel if the runoff does not re-infiltrate before reaching a stream.

Dunnian runoff, or saturation excess runoff, occurs when the subsurface storage capacity is filled (Kirkby, 1988), or when precipitation lands on the saturated areas or groundwater discharge areas, both are typically adjacent to streams. The variable source area concept (fig. 2) has been used to explain some differences between the predicted volume of storm runoff and the actual runoff because infiltration capacity at any time is only exceeded in part of a drainage basin. Betson (1964), Hewlett and Hibbert (1967), Dunne and Black (1970), and Dunne (1978) have shown that the saturated zone (variable source area) adjacent to the stream increases as the storm duration increases.

Variable Source Area Concept

Sagehen Creek Watershed, California

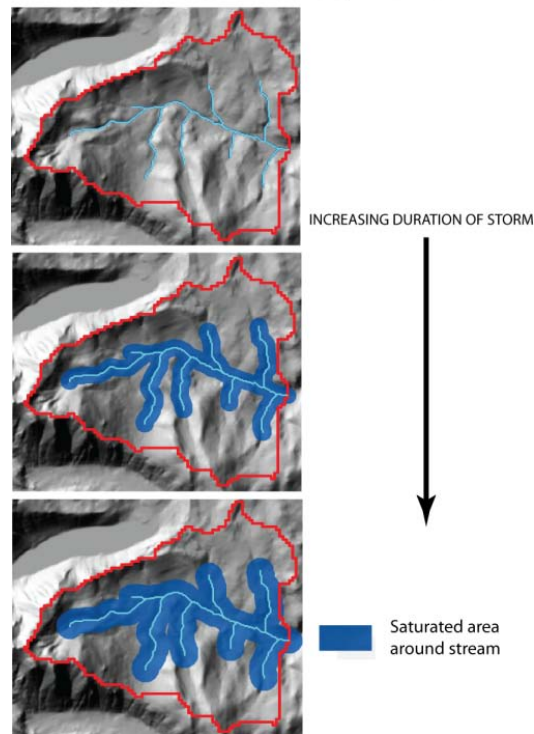


Figure 2: A hypothetical example of changes in the variable source area during a storm

The kinematic wave approximation to the Saint-Venant equations is widely used to describe open channel flow [MIKE-SHE (Refsgaard et al., 1995), HEC1 (US Army Core of Engineers, 1973), KINEROS (Woolhiser et al., 1990)]. Kinematic waves govern when internal and pressure forces are not important to the movement of the wave (i.e., the hydraulic slope is equal to the land surface slope). Typically, overland flow will be shallow, such that diffusive forces do not dominate. In the SZF Package, a quasi-2-D kinematic wave approximation is used to route overland flow according to predefined cell-by-cell pathways developed based on the slope of land surface (fig. 3).

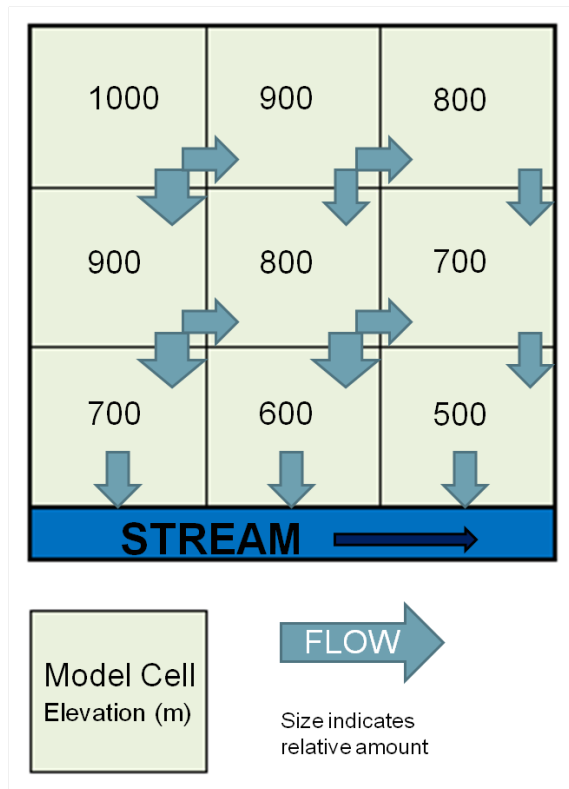


Figure 3: Example of quasi-2D cascading overland flow

DEVELOPMENT OF THE PERCHED SOIL ZONE AQUIFER

Development of a perched soil zone aquifer during a storm is illustrated in figure 4. This figure also shows macropores created by roots and faunal burrows. Development of macropores results in a higher effective hydraulic conductivity in the soil layer relative to the underlying sediment and rocks. A large fraction of infiltrating water flows into macropores, and reaches a stream through preferential-flow paths (Kirkby, 1988). Water can flow at a greater rate through macropores relative to water flowing through pores in the soil. Beven (1981) provides effective hydraulic conductivity values for soil zones in different regions that range between 1.74 m/day (Ohio) and 1,184 m/day (New Zealand).

A perched water table is a common occurrence in the soil zone because many types of restrictive layers tend to form beneath soil zones, including alluvial hardpan, unweathered rock, or a clay rich layer. The rate of rainfall or snowmelt is also an important factor that controls the formation of perched water tables, and the resulting amount of interflow. If the duration and rate of precipitation remains sufficiently high, such that the infiltration rate exceeds the percolation rate through the base of the soil zone, then interflow occurs (Dunne, 1978).

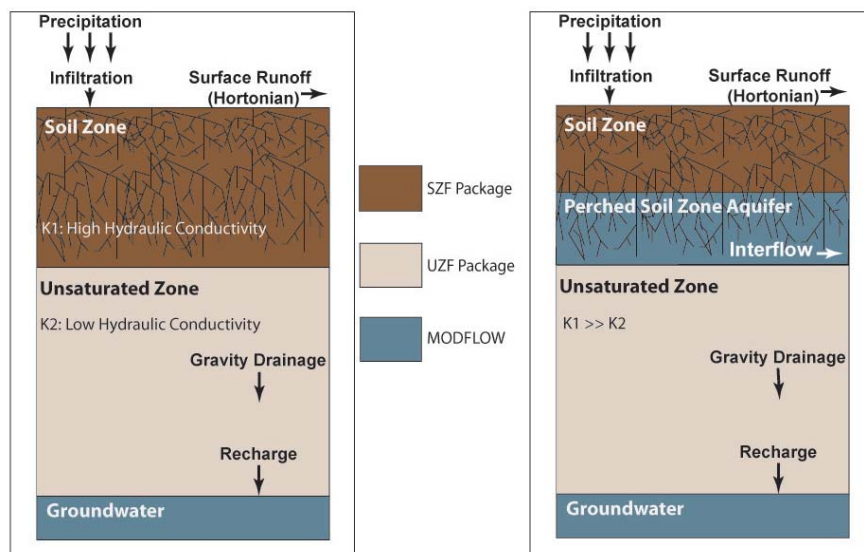


Figure 4: Conceptual model of the development of the perched soil-zone aquifer

CHAPTER 2: RESEARCH METHODS

Soil-zone flow is simulated by establishing a thin surface layer in MODFLOW with hydraulic properties representative of shallow soils (i.e., high effective hydraulic conductivity). The soil-zone layer is parameterized using available soil data including soil-zone thickness, as well as soil textural classifications that provide estimates of porosity and saturated, residual, and initial water contents, and air entry pressures. While connected macropores within the surrounding porous soil matrix may have different storage characteristics than the porous matrix itself, this model does not separately account for storage in these two domains. Hydraulic conductivity is increased to represent effective properties of both the macropores and the porous matrix in the soil zone. This enhanced permeability is what facilitates the high infiltration in the soil zone leading to interflow.

The SZF Package is combined with MODFLOW-NWT (Niswonger et al., 2011) to simulate three dimensional saturated groundwater flow, the UZF1 Package to simulate vertical unsaturated flow beneath the soil zone (Niswonger et al., 2006), and the SFR2 Package (Niswonger and Prudic, 2005) to simulate flow in streams. MODFLOW-NWT is a Newton formulation of MODFLOW-2005 (Harbaugh, 2005) that provides robust simulations of aquifers drying and rewetting, which is necessary for simulating groundwater flow in the soil zone. The SZF Package was programmed in the FORTRAN 95 programming language to interface with MODFLOW and its associated packages.

SOIL ZONE FLOW COMPUTATION

INFILTRATION AND RUNOFF

Distributed precipitation is applied to the surface layer. Precipitation applied to each MODFLOW cell is partitioned into surface runoff and infiltration using equations developed by Smith and Parlange (1978). The infiltration excess is routed as overland flow to adjacent down-slope cells using the kinematic-wave equation. Infiltration is partitioned into interflow or deep percolation based on the solution of the groundwater flow equation. Interflow is simulated using the solution to the unconfined groundwater flow equation that is solved by MODFLOW. Evapotranspiration is simulated in the soil zone as a function of soil saturation.

The infiltration rate can be computed from the following equation by Smith and Parlange (1978):

$$f_c = K_S \left[1 + \frac{\alpha}{\exp\left(\frac{\Gamma \alpha}{\Delta \theta G}\right) - 1} \right], \quad (1)$$

where f_c is the infiltration rate, K_S is the saturated hydraulic conductivity, G is the integral capillary drive across wetting front, α is the Smith-Parlange alpha dimensionless shape factor which ranges between 0.5 and 1.0, Γ is the cumulative infiltration depth, and $\Delta \theta$ is the change in water content.

The integral of capillary drive at the wetting front (G) is calculated as:

$$G = \int_{\theta_i}^{\theta_0} K(\theta) d\theta = \frac{K_s(\theta - \theta_r) \left(\frac{\theta_r - \theta}{\theta_r - \theta_s} \right)^\epsilon}{1 + \epsilon}, \quad (2)$$

where $K(\theta)$ is the unsaturated hydraulic conductivity as a function of water content (θ), θ_r is the residual water content, θ_i is the initial water content, θ_0 is the water content for the current iteration, θ_s is the saturated water content, ϵ is the Brooks-Corey shape factor.

Time of ponding is computed from the following equation (Smith-Parlange, 1978):

$$t_p = t - \left\{ \frac{(I - I_p) - \frac{\Delta\theta G K_s}{K_d} \ln \left[\frac{\exp\left(\frac{\alpha I'}{\Delta\theta G}\right) - 1 + \gamma}{\exp\left(\frac{I_p' \alpha}{\Delta\theta G}\right) - 1 + \gamma} \right]}{K_s(1 - \alpha)} \right\} \quad (3)$$

where t is time, t_p is time of ponding, α is Smith-Parlange alpha, I is the cumulative infiltration, I_p is the ponded infiltration, K_d is the difference between $K(\theta_i)$ and $K(\theta_2)$, θ_2 is water content at time= t_2 , γ is a dimensionless shape factor equal to $\alpha * K_s / K_d$, and I_p' has the following relationship;

$$I_p' = (I_p) - K_i * t_p.$$

The cumulative infiltration at time of ponding is computed as the rainfall rate multiplied by the time of ponding.

Soil moisture in the soil zone between rainfall events is redistributed using the following equation by Smith-Corradini (1993). This ensures that the next time-of-ponding calculation takes into account updated soil moisture condition. The equation can be written as:

$$\frac{d\theta_0}{dt} = \frac{\Delta\theta_{i0}}{I'_z} \left[R_h - K_i - \left(K(\theta_0) + \frac{\beta p K_s \Delta\theta_{i0} G}{I'_z} \right) \right], \quad (4)$$

where θ_0 is water content at land surface, $\Delta\theta_{i0}$ is the difference between θ_0 at land surface and θ_i , I'_z is the infiltrated water during hiatus, R_h is the rainfall rate during hiatus, K_i is the hydraulic conductivity at θ_i ($K(\theta_i)$), β is the Smith-Parlange Beta dimensionless shape factor which ranges between $\Pi/4$ and 1.0, p is the dimensionless wetting front shape factor which ranges from 1.0-2.0, K_s is the saturated hydraulic conductivity, and G is the integral capillary drive across the wetting front.

OVERLAND FLOW

Surface runoff for each MODFLOW finite-difference cell is calculated and routed to the nearest channel using a Quasi 2-D kinematic-wave approximation. Discrete 1-D pathways are defined based on relative elevations defined by a digital elevation model (DEM). Four pathways can be simulated for each grid cell, similar to a 2-D finite-difference formulation. Kinematic waves then route the overland flow downhill along these paths to stream channels. Runoff that flows from one model cell to multiple

downslope cells is partitioned according to the relative slope among the downslope cells. If surface runoff flows to a downstream cell that has more infiltration capacity, then all or a portion of the surface runoff re-infiltrates.

Hortonian runoff is calculated using the time-of-ponding relationship between the precipitation rate and the infiltration rate. If precipitation is applied to a MODFLOW cell that is already saturated, Dunnian saturation excess runoff will be generated. If interflow moving through the soil zone encounters an area of saturation then return flow is routed to the surface and overland flow continues to the next model cell. If precipitation is applied to an impervious surface, such as pavement, the precipitation is automatically added to overland flow. The 1-D kinematic wave approximation (Lighthill and Whitham , 1955) is used to route overland flow for each downstream connection as:

$$\frac{1}{c} \frac{\partial q}{\partial t} + \frac{\partial q}{\partial x} = \frac{Q(x,t)}{l} \quad , \quad (5)$$

where C is Chezy's coefficient, x is 1-D flow path direction, l is cell length along columns or row (depends on flow direction out of each cell), $Q(x, t)$ is volume flux, and q is the flow velocity. Chezy's relation is used to calculate mean flow velocity as:

$$q = C\sqrt{RS_0} \quad , \quad (6)$$

where R is the hydraulic radius, S_0 is land surface slope, and C is Chezy's coefficient.

Equation 6 is solved using a 4-point finite-difference scheme (Preissmann, 1961)

INTERFLOW AND GROUNDWATER FLOW

Interflow and vertical flow beneath the base of the soil zone is calculated according to the groundwater-flow equation solved by MODFLOW-NWT. If the groundwater head is at or above the base of the soil zone, then the soil zone cell is treated as an aquifer. If the groundwater head is below the base of the soil zone, then water is routed through the deeper unsaturated zone by the UZF1 Package. Interflow through the soil zone and regional groundwater flow is modeled using the three-dimensional groundwater-flow equation (Harbaugh, 2005; Niswonger et al., 2011):

$$\frac{d}{dx}\left(K_{xx}\frac{dh}{dx}\right) + \frac{d}{dy}\left(K_{yy}\frac{dh}{dy}\right) + \frac{d}{dz}\left(K_{zz}\frac{dh}{dz}\right) + W = S_s\frac{dh}{dt}, \quad (7)$$

where K_{xx} , K_{yy} , K_{zz} are hydraulic conductivity in the x, y, and z directions, h is the groundwater head, W represents any sources or sinks of water, and S_s is specific storage. Volume averaging of equation (7) results in the conservative form of the groundwater flow equation (Harbaugh, 2005; Markstrom et al., 2008):

$$\Sigma Q_L = S_s\frac{dh}{dt}V \quad (8)$$

where Q_L is the volumetric flow rate into a cell (negative value is a volumetric flow out of cell, in cubic length per time), and V is the volume of the cell in cubic length.

IMPLEMENTATION OF INTERFLOW BOUNDARY INTO MODFLOW

Interflow to streams is implemented through a head-dependent seepage face in each cell that contains a stream. Interflow is calculated on the basis of the head in the model cell, the stream elevation, and the stream stage. The interflow calculation considers the feedback between stage, streamflow, and interflow; which provides numerical stability.

The interflow boundary is implemented into the MODFLOW-NWT as a nonlinear stress. Niswonger et al. (2011) provide a detailed explanation of how a nonlinear stress is added to MODFLOW-NWT. The function added to MODFLOW-NWT for the calculation of interflow is

$$f(h) = K_{if} S_n T (B - h), \quad (9)$$

where, K_{if} is the interflow conductance term, $S_n * T$ is a function which smoothes conductance as a function of saturated thickness in the cell; B is the bottom elevation of the soil-zone cell if the stream stage is not above the soil-zone base, or the stream stage elevation, if it is above the soil zone base; h is the head in the soil zone in the model cell which contains the stream. The interflow conductance term is

$$K_{if} = Hk \frac{S * h}{\frac{1}{4} * (delr + delc)}, \quad (10)$$

where Hk is the horizontal hydraulic conductivity of the soil zone, S is two times the stream length of the reach in the model cell for which interflow is calculated, h is the head in the soil-zone model cell that contains the seepage face to the stream, $delr$ is the model cell dimension in the row direction, and $delc$ is the model cell dimension in the column direction.

Twice the stream length is multiplied by h to calculate the facial area for interflow because the stream length is the distance over which interflow will occur on each side of the stream channel. Half of the average of the cell dimensions, $delr$ and $delc$ represents the distance over which the hydraulic gradient is calculated between the edge of the model cell and the model node. The average of $delr$ and $delc$ is used because streams meander within model cells, and are sub grid-scale features for which obtaining an exact distance from the model node to each point in the stream reach is not available.

The Newton formulation is

$$\frac{d}{dh}(f(h))\Delta h = -f(h) ,$$

where in this example, $f(h)$ is equation (9), and $\Delta h = h_k - h_{k-1}$ where, the subscripts k and $k - 1$ designate values of h for nonlinear iterations k and $k - 1$. Implementing the Newton formulation, equation (9) becomes

$$\frac{\partial}{\partial h} [K_{if} Sn T (B - h)] \Delta h = -[K_{if} Sn T (B - h)]. \quad (11)$$

Evaluating the derivative and distributing Δh results in

$$\left[K_{if} \frac{\partial S_n}{\partial h} (B - h_{k-1}) - K_{if} S_n T \right] h_k = \quad (12)$$

$$\left[K_{if} \frac{\partial S_n}{\partial h} (B - h_{k-1}) - K_{if} S_n T \right] h_{k-1} - \left[K_{if} S_n T (B - h_{k-1}) \right],$$

To implement the interflow into MODFLOW-NWT, equation (12) is added to the discretized groundwater flow equations.

TEST SIMULATION

A hypothetical model, similar to the model used by Vauclin et al. (1979), was used to compare results from a MODFLOW-SZF simulation with a two dimensional form of Richards' equation (VS2DT; Healy, 1990). The test simulation has two hydrogeologic units (table 1) represented in two model layers consisting of a highly permeable layer, which represents the soil zone, overlying a lower permeable clay layer, which represents the regional aquifer and unsaturated zone (fig. 5). Hydraulic conductivity of the soil zone ranged from a maximum value of 250 m/d to a minimum of 0.05 m/day in several simulations (table 2). The vertical and horizontal hydraulic conductivity were parameterized to represent the extensive preferential flow path networks typically present in the soil zone.

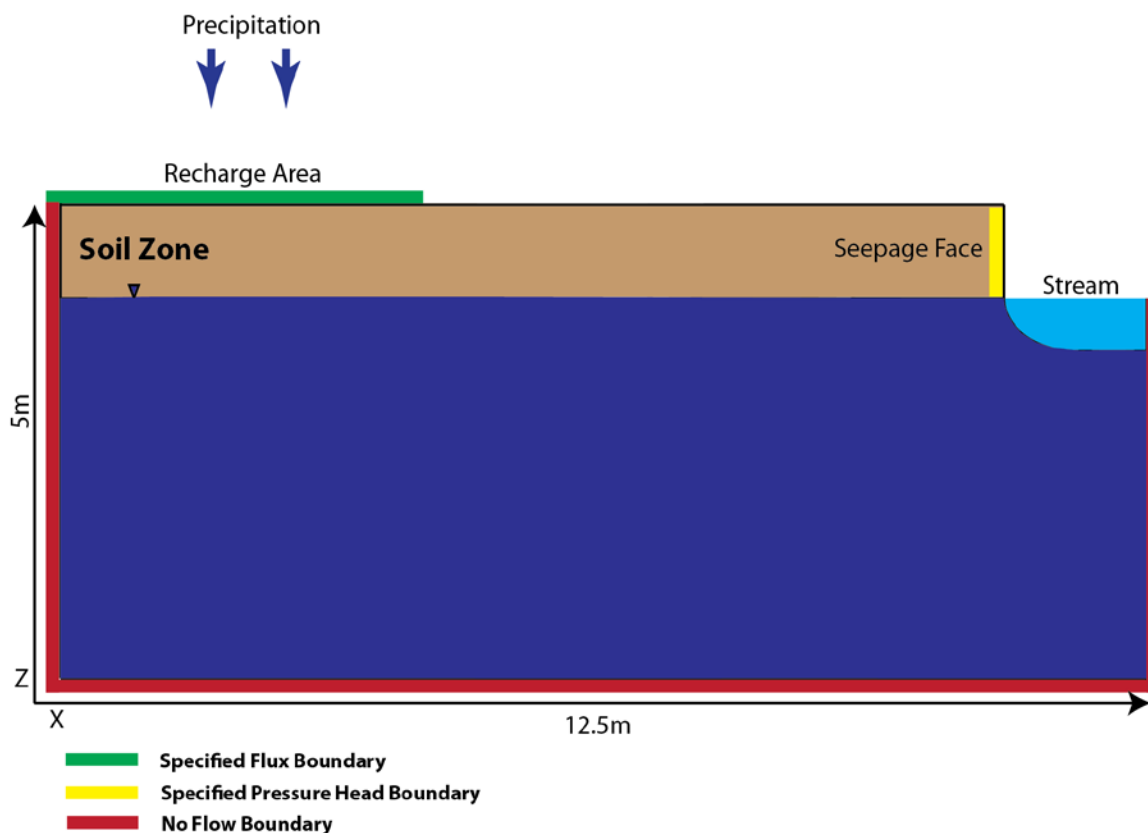


Figure 5: Model domain for test simulations comparing results from MODFLOW-SZF with VS2DT.

The cross-sectional model was 12.5 m in the x-direction, 5 m in the z-direction, and 1 m in the y-direction. At the distal edge of the cross section, interflow occurs as a seepage face to the stream channel. The same hydraulic properties were used for both simulations (table 1) using values typical for each hydrogeologic unit. The water table was set equal to the elevation at the base of the soil zone. Constant precipitation was applied on the soil-zone surface over a 5 m² area adjacent to the model boundary, opposite the interflow boundary for the entire 20-day simulation. In the model comparison, 0.25, 0.5, and 1.0 day timesteps were evaluated in MODFLOW-SZF and in

VS2DT adaptive time steps were used which range from a minimum size of 0.005 days to a maximum of 1 day.

Table 1: Input parameters for SZF and VS2DT in simulation 1

Parameter	Layer 1	Layer 2
Horizontal Hydraulic Conductivity (m/d)	250	0.002
Vertical Hydraulic Conductivity (m/d)	0.5	0.002
Specific Storage	1.0E -5	1.0E -5
Specific Yield	0.35	0.15
Porosity	0.45	0.45
Residual Moisture Content	0.10	0.30
Air Entry Pressure	-0.05	-0.05
Lambda	2	2
Epsilon	4	4

The focus for this thesis will be on model results from four simulations using the SZF Package combined with MODFLOW-NWT (referred to as MODFLOW-SZF). These simulations test the effects of varying: 1) grid dimensions, 2) hydraulic conductivity values, and 3) rainfall rates. Simulation 1 used the input parameters listed in table 1 to evaluate the effectiveness of using the groundwater-flow equation to represent interflow in MODFLOW-SZF by comparing with Richards' equation in VS2DT. Additionally, the effects of grid-cell size was varied by testing grid-cell dimensions ranging between 0.05 and 0.5m in VS2DT, and horizontal cell dimensions ranging between 0.25 and 5 m with vertical discretization ranging between 0.25 to 1 m in MODFLOW-SZF. Testing the effects of grid-cell dimensions in VS2DT provided a means of establishing a “correct” solution for comparison to MODFLOW-SZF for the hypothetical interflow problem. The grid-cell dimensions did not have an effect on the VS2DT simulation results for cell dimensions less than 0.1 m. Accordingly, the VS2DT

model with a grid cell dimension equal to 0.1 m or smaller was used to represent the “correct” solution for evaluating the accuracy of MODFLOW-SZF. Three additional simulations were run to examine sensitivity to vertical and horizontal hydraulic conductivity and applied rainfall rate. VS2DT does not simulate runoff; therefore, runoff was not a part of the test simulation comparison. Table 2 shows a summary of the parameters that were varied in all of the simulations..

Table 2: Summary of parameters varied in simulations [HK: horizontal hydraulic conductivity; VK: vertical hydraulic conductivity]

Simulation	HK (m/day)	VK (m/day)	Precipitation rate (m/day)	Total Infiltration (m ³)	Time (Days)
1	250	0.50	0.04	4.00	20
2	25	0.05	0.04	4.00	20
3	250	0.50	0.08	8.00	20
4	250	0.50	0.12	12.00	20

The effects of grid-cell dimensions on the amount of interflow simulated by MODFLOW-SZF was evaluated by comparing results from the three different grid-cell configurations. Three different grid-cell configurations were used in MODFLOW-SZF to test the effects of grid-cell dimensions on the interflow solution. Configuration 1 was a coarse resolution model with uniform 5 m spacing in the x-direction for both layers. Layer 1 was 1 m thick and layer 2 was 4 m thick. Configurations 2 and 3 represent finer resolution models. Configuration 2 had uniform spacing of 0.25 m in the x-direction for both layers with the same vertical resolution as configuration 1. Configuration 3 had the finest resolution with a grid-cell dimension of 0.25 m, in the x and z directions in both layers.

Four different grid-cell configurations were used in VS2DT to test the effects of grid-cell dimensions on the interflow solution. Uniform grid-cell dimensions in the x and z directions were used in all VS2DT simulations equal to 0.05 m, 0.10 m, 0.25 m, and 0.50 m, configurations 1, 2, 3 and 4 respectively. MODFLOW-SZF and VS2DT solutions were compared based on timing of the vertical and horizontal wetting fronts, flow at the seepage face, and development of a water table.

Sensitivity analyses were done (referred to as simulations 2, 3, and 4) based on MODFLOW-SZF configuration 1 and VS2DT configuration 2. Sensitivity analyses were used to evaluate the effect of various precipitation rates and hydraulic-conductivity values on the timing and amount of interflow. The simulations were designed to analyze and compare the timing of interflow and percent difference of total interflow to the stream when varying hydraulic conductivity (simulation 2) and precipitation (simulations 3 and 4).

Model sensitivity to varying hydraulic conductivity was evaluated by comparing results between simulations 1 and 2, in which the ratio between vertical and horizontal hydraulic conductivity remained constant at 0.002 and horizontal and vertical hydraulic conductivity of the soil zone was decreased by an order of magnitude (table 2).

Sensitivity to precipitation rates was evaluated in simulations 1, 3, and 4, where applied precipitation rates increased from 0.04 m/d in simulation 1, to 0.08 m/d in simulation 2, and to 0.12 m/d in simulation 3.

In all of the test simulations, the interflow percent difference was calculated as the difference in total interflow between models divided by the average interflow. Root-

mean-squared error (RMSE) was calculated as the square root of the average of the squared differences between model simulations.

CHAPTER 3: RESULTS AND DISCUSSION

DISCRETIZATION

Figures 6 and 7 show the effects of different grid-cell configurations. All MODFLOW-SZF grid-cell configurations had almost identical interflow breakthrough curves (fig. 6). This result indicates that the MODFLOW-SZF interflow solution is not sensitive to varying grid-cell dimensions over the range evaluated in these simulations. Additionally, time discretization had no effect on the MODFLOW solutions for time steps equal to and smaller than one day.

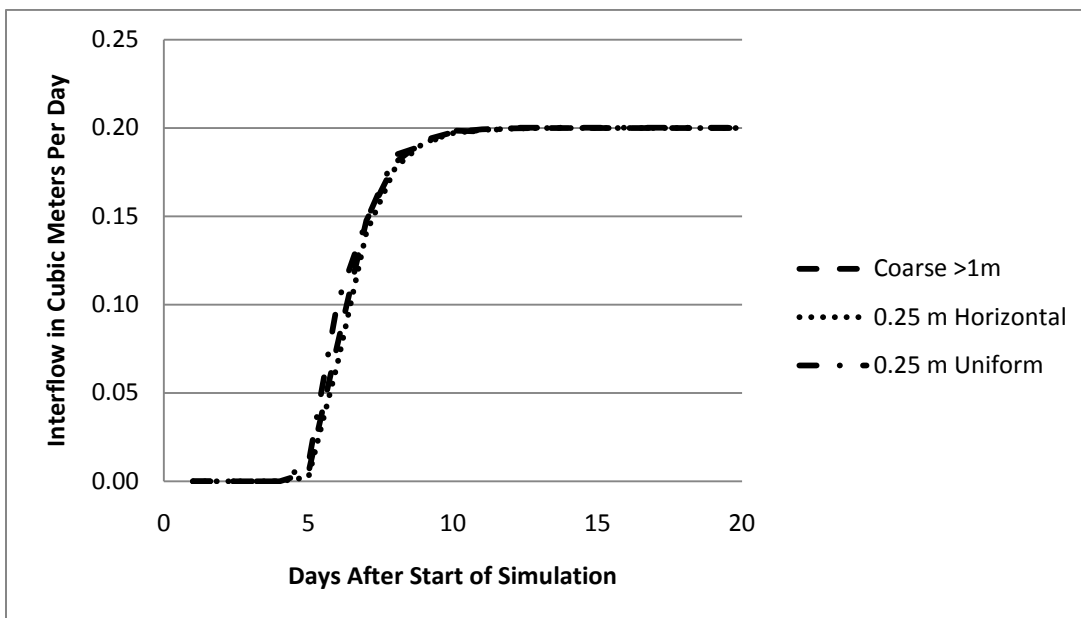


Figure 6: Effects of varying grid-cell dimensions on interflow simulated by MODFLOW-SZF

The interflow percent difference and RMSE for the smaller grid-cell configurations (0.25 x 1 m, and 0.25 x 0.25 cm) was calculated relative to the largest grid-cell configuration (5 x 1 m) because evaluation of MODFLOW-SZF interflow solutions at coarser grid resolution is the focus of this research. The interflow percent differences were low for all MODFLOW-SZF configurations, ranging between 0.92 and 1.10%. RMSE values were also low, ranging between 0.003 and 0.01.

Table 3: Summary of effects of varying grid-cell dimensions on interflow simulated by MODFLOW-SZF

Configuration	Discretization	Total Interflow (m ³)	Interflow % Difference	RMSE
1	5 x 1 m	2.80	-	-
2	0.25 X 1 m	2.77	0.92%	0.003
3	0.25 x 0.25 m	2.83	1.10%	0.01

The effects of varying grid-cell dimensions were much more pronounced in VS2DT (fig. 7 and table 4). Grid-cell dimensions affected both the timing and the shape of the interflow-breakthrough curves.. Simulations that used uniform grid-cell dimensions of 0.25 m and 0.5 m resulted in a 5 and 10-day delay, respectively, in the onset of interflow relative to the smaller grid-cell configurations (0.1 and 0.05 m; fig. 7). Reducing the grid-cell dimensions lower than 0.1 m had a negligible effect on the interflow simulated by VS2DT.

The interflow percent difference and RMSE for each VS2DT grid-cell configuration (uniform 0.1, 0.25, and 0.50 m) were calculated relative to the model with the smallest grid-cell dimension (0.05 x 0.05 m). The interflow percent difference ranged

between 0.02% and 82.6%; the RMSE ranged between 0.008 and 0.12. As expected, the error in the interflow was larger for larger grid-cell dimensions.

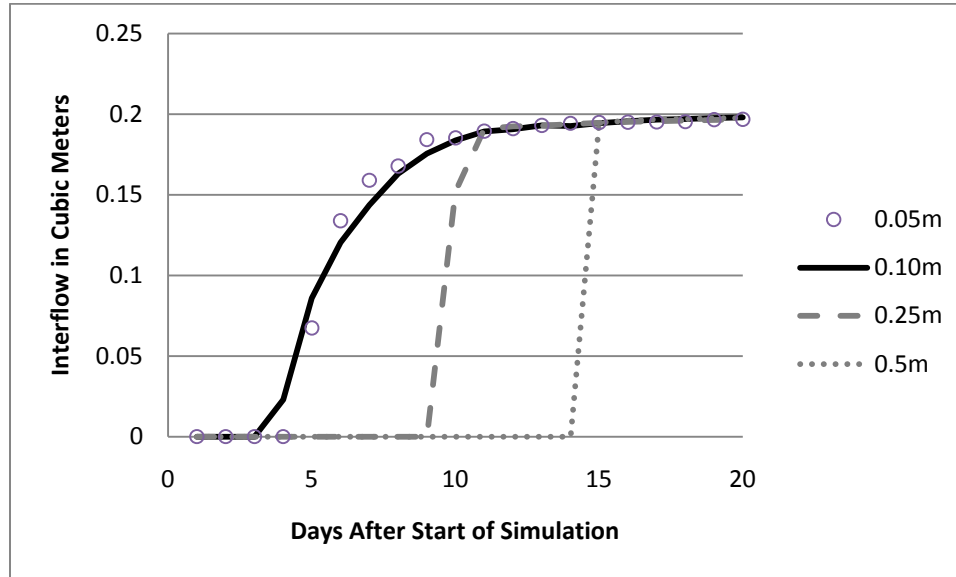


Figure 7: Effects of grid-cell dimensions on interflow simulated by VS2DT

Table 4: Summary of effects of grid-cell dimensions on interflow simulated by VS2DT

Configuration	Discretization	Total Interflow (m ³)	Interflow % Difference	RMSE
1	0.05 x 0.05 m	2.83	-	-
2	0.1 x 0.1 m	2.84	0.02%	0.008
3	0.25 x 0.25 m	2.09	30.09%	0.074
4	0.50 x 0.50 m	1.18	82.60%	0.12

Grid-cell dimensions of 0.05 and 0.1 m provided the most stable VS2DT solutions, and these results are in agreement with results presented by Jensen and Montglou (1992). The 0.1 m discretization was the largest cell size that preserves a stable

solution. For these reasons, the 0.1 m uniform grid VS2DT model is the model that was compared with MODFLOW-SZF for all of the sensitivity analyses.

COMPARISON OF MODFLOW-SZF AND VS2DT

The effects of grid-cell dimensions on model computation time were compared between MODFLOW-SZF and VS2DT (table 5). VS2DT with a 0.1 m grid-cell dimension required a total of 6,250 grid cells and completed computation in 40 seconds. In contrast, MODFLOW-SZF required only 6 grid cells to obtain an accurate solution and completed computation in 0.11 seconds

Table 5: Summary of model computation times for MODFLOW-SZF and VS2DT

	MODFLOW		VS2DT	
	1	3	2	3
Configuration				
Discretization	5 x 1 m	0.25 x 0.25 m	0.1 x 0.1 m	0.25 x 0.25 m
# Model Nodes	6	1020	6,250	1,000
Computation Time (seconds)	0.11	22.5	40	10

Figure 8 shows the comparison between MODFLOW-SZF configuration 1 and VS2DT configuration 2. For simulation 1, interflow simulated by VS2DT begins one day earlier and initially has a steeper curve than interflow simulated by MODFLOW-SZF; however, interflow from VS2DT approaches steady state slower (fig. 8). The total fluxes and storage agree well with the VS2DT model over the duration of the simulation with a low RMSE of 0.022 between models.

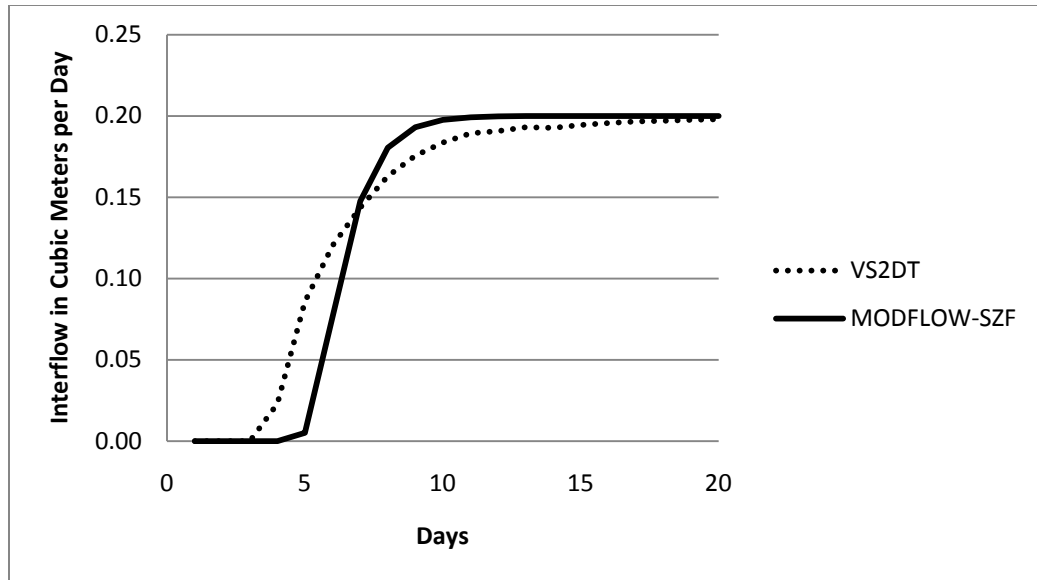


Figure 8: Comparison of interflow from simulation 1 using the best fit discretizations of MODFLOW-SZF and VS2DT with horizontal hydraulic conductivity of 250 m/day and vertical hydraulic conductivity of 0.50 m/day for the soil zone.

WETTING FRONT

Arrival of the wetting front at a depth of 1 m (i.e. the base of the soil zone) occurred at nearly the same time for the MODFLOW-SZF and VS2DT simulations (fig. 9). However, the wetting front simulated by VS2DT is more diffuse than that simulated by MODFLOW-SZF, owing to the effects of capillary pressure gradients at the wetting front. The UZF1 Package is used to route wetting fronts vertically in MODFLOW-SZF, and the UZF1 Package does not consider capillary pressure gradients. VS2DT simulations showed that the hydraulic conductivity averaging method used in VS2DT also affected the shape of the wetting front.

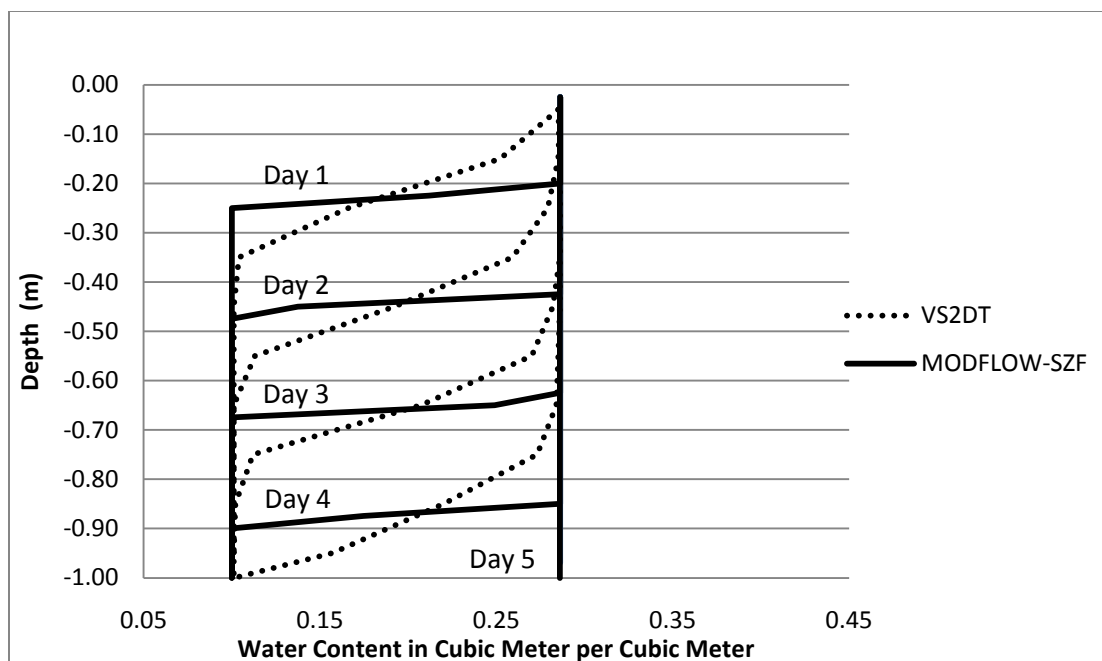


Figure 9: Comparison of vertical wetting fronts from simulation 1 between MODFLOW-SZF and VS2DT

WATER TABLE

Changes in the water table caused by precipitation and infiltration on the left side of the model domain (fig. 5) were compared between MODFLOW-SZF and VS2DT after 5 and 7 days and at steady state (fig. 10). The 0.25-m horizontal grid spacing (configuration 2) for MODFLOW-SZF was used because it provided a better horizontal resolution of the water table and reduced the need to interpolate water table elevations between largely-spaced model nodes. Comparison with MODFLOW-SZF configuration 2 is acceptable in this case because interflow results from MODFLOW-SZF were independent of the horizontal grid spacing (fig. 6).

The water table in VS2DT was calculated based upon pressure head at the lowest soil-zone model node and the difference between model node elevation and the base of

the soil zone. The water table was assumed equal to the node elevation when the pressure head at the node was equal to zero. If the elevation of the zero pressure condition occurred between nodes, then the location of the water table was interpolated between the nodes.

The water table was higher in MODFLOW-SZF because of the differences in how unsaturated storage is handled in each model. MODFLOW-SZF resolves the water table at a sub-grid scale using the vertical equilibrium assumption, which assumes that all water occupies the bottom of the cell, whereas in Richards' equation the volume of water contained in a cell is averaged over the volume of the cell. Capillarity also affects the water table height. In VS2DT, water is pulled from the saturated zone into the capillary fringe, whereas in MODFLOW water is not stored in a capillary fringe. The wetting front reaches the base of the soil zone at five days in both models and the steady state water table occurs at ten days in each model.

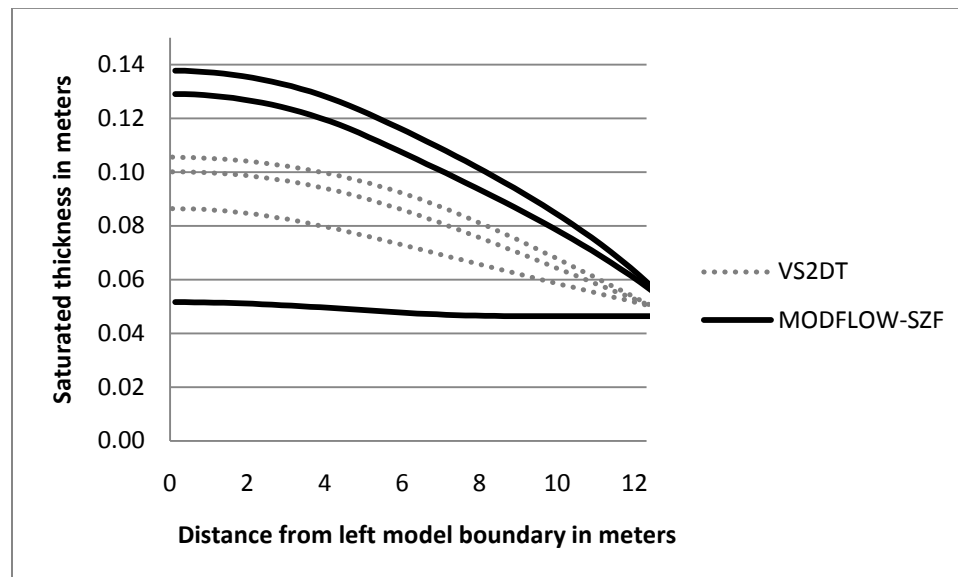


Figure 10: Comparison of rising water tables during recharge for 0.25 m spacing in MODFLOW-SZF and 0.1 m spacing in VS2DT and for simulation 1 at 5 days, 7 days, and steady state.

SENSITIVITY ANALYSES

HYDRAULIC CONDUCTIVITY

The effect of varying the hydraulic conductivity on interflow was evaluated by decreasing the horizontal hydraulic conductivity an order of magnitude from 250 m/d to 25 m/d while keeping the vertical to horizontal hydraulic conductivity ratio (0.002) the same (table 2). For this simulations with a lower hydraulic conductivity, VS2DT simulated the onset of interflow after 6 days as compared with 7 days for MODFLOW-SZF, although interflow does not increase as rapidly for VS2DT (fig. 11). Interflow curves between MODFLOW-SZF and VS2DT are similar with a RMSE of 0.014.

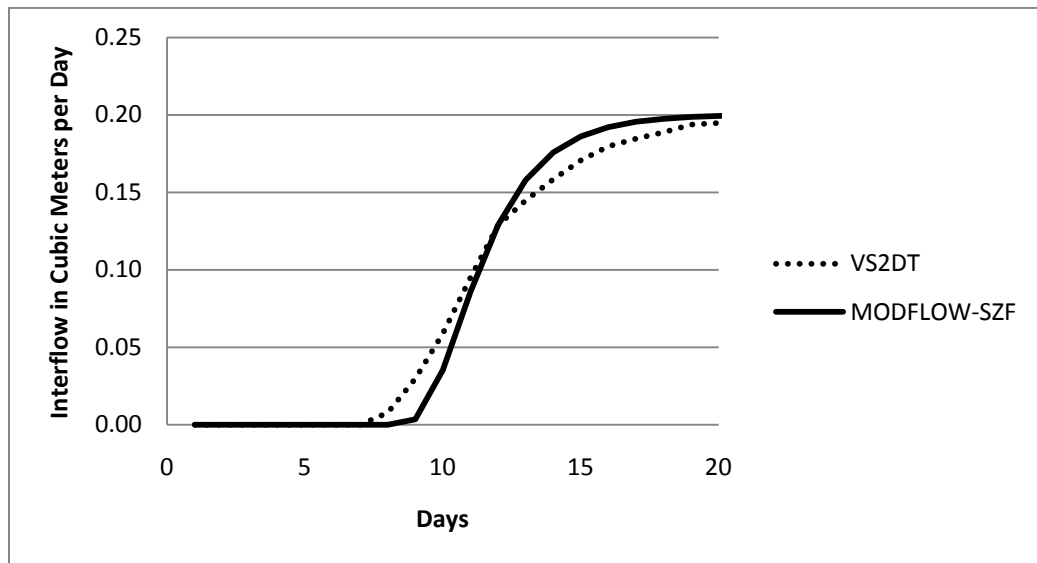


Figure 11: Comparison of interflow simulated by MODFLOW-SZF and VS2DT with horizontal hydraulic conductivity of 25 m/day and vertical hydraulic conductivity of 0.05 m/day

PRECIPITATION RATE

Figures 12 and 13 show model sensitivity to varying precipitation rates. The applied precipitation rate was 0.08 m/day in simulation 3, and the applied precipitation rate was 0.12 m/day in simulation 4 (table 2). Interflow occurred one day earlier in VS2DT compared with MODFLOW-SZF in simulations 3 and 4, and both VS2DT simulations reached steady state about 1 day after MODFLOW-SZF reached steady state. The shapes of the interflow curves are similar in simulations 3 and 4 with low RMSE of 0.045 and 0.052, respectively.

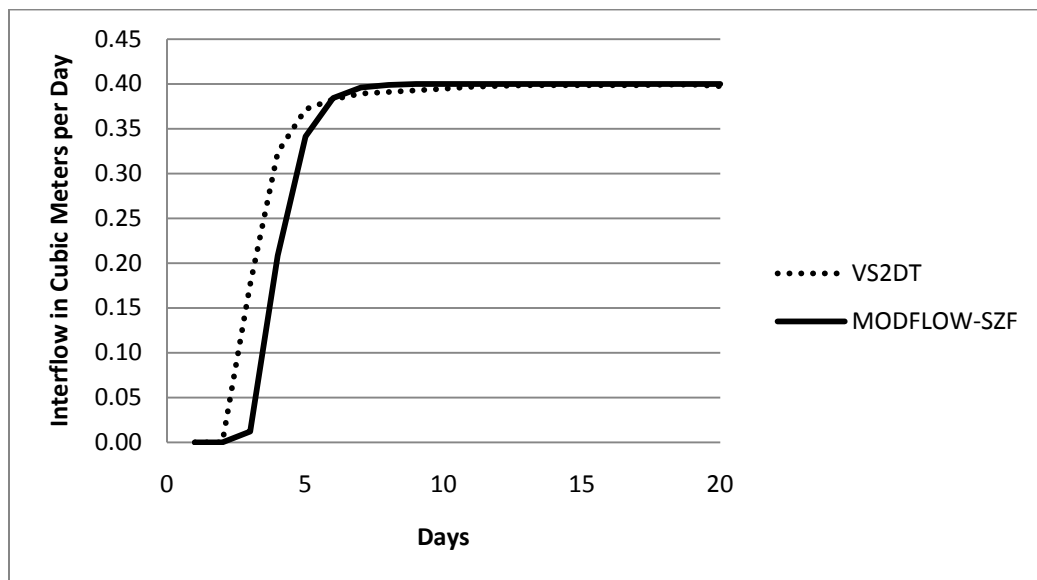


Figure 12: Comparison of interflow simulated by MODFLOW-SZF and VS2DT for an applied precipitation rate of 0.08 m/day

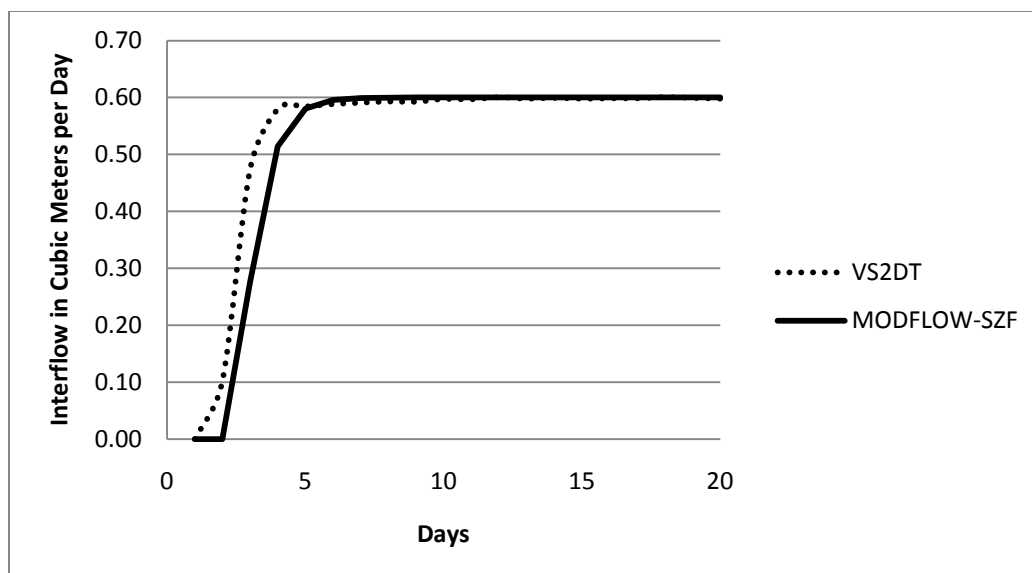


Figure 13: Comparison of interflow simulated by MODFLOW-SZF and VS2DT for an applied precipitation rate of 0.12 m/day

Table 6 summarizes the total interflow, percent differences, and RMSE values for the different simulations that were analyzed in this sensitivity analysis. The MODFLOW-SZF and VS2DT simulations had similar results for soil storage and interflow. The difference in total fluxes can be attributed to the timing delay of steady-state flux through the seepage face. The percent differences between MODFLOW-SZF and VS2DT were all less than 1.5% in all of the simulations with small RMSE ranging from 0.014-0.052.

Table 6: Summary of results of sensitivity analyses [storage and total interflow values in cubic meters]

Simulation	MODFLOW		VS2DT		Interflow % difference	RMSE
	Storage	Total Interflow	Storage	Total Interflow (m ³)		
1	1.20	2.80	1.22	2.78	0.72%	0.022
2	2.32	1.68	2.32	1.68	0.00%	0.014
3	1.46	6.53	1.38	6.62	-1.37%	0.045
4	1.63	10.37	1.57	10.43	-0.58%	0.052

IMPLICATIONS

The errors associated with using the groundwater flow equation to represent flow instead of Richards' equation were much lower (table 6) than the errors caused by grid effects (table 4), across a range of hydraulic conductivities and applied precipitation rates. In order to obtain a stable solution for Richards' equation, the model resolution had to be 0.1 m or less. When you consider that regional models encompass hundreds to thousands of square kilometers, using a grid resolution of 0.1 m quickly becomes intractable, especially if results are desired in less than real time. In basin to regional scale applications, the spatial scale will be even greater than was used in this comparison, indicating that the groundwater flow equation is likely to be more accurate at representing interflow through the soil-zone when applied to scales greater than an experimental plot.

Recent research in this area of hydrology has been focused developing efficient numerical methods for solving Richards' equation in the context of integrated basin-scale models. These results suggest that more evaluation of simple governing equations is needed to determine how to best to represent variably saturated flow processes at the basin scale.

CHAPTER 4: SUMMARY AND CONCLUSIONS

Models can be improved by considering subsurface conditions in the soil zone because infiltration and runoff are functions of soil-zone saturation. The SZF Package uses Smith-Parlange infiltration equations, quasi 2-D kinematic wave overland flow routing, and the groundwater flow equation to simulate infiltration, runoff, and interflow because these equations are generally not as sensitive to grid-cell dimensions as Richards' equation.

- The Soil-Zone Flow Package for MODFLOW effectively simulates infiltration and interflow through the soil zone; verified through comparison to VS2DT including: water table development, timing of the wetting front from the land surface to the base of the soil zone, timing of interflow to the stream, soil-zone storage, total infiltration, and total interflow for a variety of applied precipitation rates and hydraulic conductivities. Timing of vertical and horizontal wetting fronts in both models was consistently less than 1 day of each other.
- The MODFLOW-SZF simulations achieved a reasonable infiltration and interflow solution with only 6 grid cells and 0.11 seconds of computation time, requiring a fraction of the cells and computation time required for VS2DT.
- The MODFLOW-SZF interflow solution was independent of discretization, whereas the VS2DT interflow solution was strongly sensitive to the discretization with percent differences between the 0.05-m and corresponding 0.1-m, 0.25-m

and 0.5-m discretizations equal to 0.02%, 0.30%, and 82.6% respectively and RMSE between 0.008 and 0.012.

- MODFLOW-SZF simulations correlated well with the VS2DT simulations, with simulation total percent differences of 1.5% or less and RMSE between 0.014 and 0.052 in all of the hydraulic conductivity and precipitation sensitivity simulations.
- The errors associated with discretization in Richards equation (up to 82.6% total interflow percent difference, RMSE up to 0.12) far outweigh the errors associated with using the groundwater flow equation to represent soil-zone flow (<1.5% total interflow percent difference, RMSE <0.052).

The SZF Package was developed for MODFLOW to incorporate soil-zone processes at spatial scales useful for basin-scale models. Results suggest that the Smith-Parlange analytical solution to Richards' equation and the groundwater flow equation can be used to accurately represent soil-zone processes at a variety of spatial scales, precipitation rates, and hydraulic conductivities independent of discretization. The SZF Package effectively simulates infiltration and interflow at a coarser grid resolution than is usually required to obtain a stable solution to Richards' equation, indicating that it will be an effective framework for modeling soil zone processes at the regional scale.

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APPENDIX A: SZF FORTRAN CODE

```

C-----SUBROUTINE GWF2PSZ1AR
C
  SUBROUTINE GWF2PSZ1AR(IN,IUNITUZF,IUNITSFR,IGRID)
C
  USE GLOBAL,          ONLY: NCOL, NROW, NLAY, IOUT, ITRSS, ISSFLG,
+                        DELR, DELC, IBOUND, BOTM, HNEW
  USE GWFPSZMODULE
  USE GLOBAL,          ONLY: ITMUNI, LENUNI
  IMPLICIT NONE
C
  -----
C  SPECIFICATIONS:
C  -----
C  ARGUMENTS
C  -----
  INTEGER IN, IGRID, IUNITUZF, IUNITSFR
  CHARACTER*24 ANAME(12)
  DATA ANAME(1)/' PERCH VERT. HYD. COND. '/
  DATA ANAME(2)/' SZ AIR ENTRY PRESSURE  '/
  DATA ANAME(3)/' SZ SPECIFIC RETENTION  '/
  DATA ANAME(4)/' SZ POROSITY             '/
  DATA ANAME(5)/' SZ B-K EPSILON         '/
  DATA ANAME(6)/' SZ SAT.VERT. HYD. COND. '/
  DATA ANAME(7)/' SZ HYD.COND. ALONG COL. '/
  DATA ANAME(8)/' SZ HYD.COND. ALONG ROWS'/
  DATA ANAME(9)/' SZ PARLANGE ALPHA      '/
  DATA ANAME(10)/' SZ INIT. WATER CONTENT '/
  DATA ANAME(11)/' CASCADE SURFACE       '/
  DATA ANAME(12)/' CHEZY"S COEFFICIENT   '/
C
  -----
C  LOCAL VARIABLES
C  -----
  INTEGER IUZM, LLOC, ISTART, ISTOP, ICONFLAG, NUMCHECK
  INTEGER IC, IR, L
  CHARACTER*200 LINE
  REAL R
C
  -----
  ALLOCATE (PERCHFLG, IPSZCB1, IPSZCB2, IFILLFLG, IUNITPSZ )
  IUNITPSZ=IN
C1-----IDENTIFY PACKAGE AND INITIALIZE.
C2-----PERCHFLG.GT. ZERO VKS FROM UZF/PERCHFLG.LE.0 VKS FROM PSZ
  WRITE (IOUT, 9001) IN
  9001 FORMAT (1X, /' PSZ1 -- PERCHED FLOW AND SOIL ZONE PACKAGE, ',
+            'VERSION 1, 10/03/2005', /, 9X, 'INPUT READ FROM UNIT', I3)
  CALL URDCOM(IN, IOUT, LINE)
  LLOC = 1
  CALL URWORD(LINE, LLOC, ISTART, ISTOP, 2, PERCHFLG, R, IOUT, IN)
  IF ( PERCHFLG.GT.0 .AND. IUNITUZF.EQ.0 ) THEN
    WRITE (IOUT, 9002)
  9002 FORMAT (/' VERTICAL HYDRAULIC CONDUCTIVITY ',
+            'OF THE PERCHING LAYER IS TO BE READ FROM UZF '
+            //, ' AND UZF IS INACTIVE -- SETTING PSZ PACKAGE TO
+            'INACTIVE'///)

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```

        IN = 0
        RETURN
    ELSE IF ( PERCHFLG.LE.0 ) THEN
        WRITE (IOUT, 9003)IN
9003  FORMAT (//' READING VERTICAL HYDRAULIC CONDUCTIVITY ',
+          //'OF THE PERCHING LAYER FROM UNIT ',I5//)
        ELSE
            WRITE (IOUT, 9004)
9004  FORMAT (//' THE VERTICAL HYDRAULIC CONDUCTIVITY (K) ',
+          //'OF THE PERCHING LAYER IS BEING SET ',
+          'EQUAL TO THE VERTICAL K FROM UZF'//)
        END IF
        CALL URWORD(LINE, LLOC, ISTART, ISTOP, 2, IPSZCB1, R, IOUT, IN)
        CALL URWORD(LINE, LLOC, ISTART, ISTOP, 2, IPSZCB2, R, IOUT, IN)
        CALL URWORD(LINE, LLOC, ISTART, ISTOP, 2, IFILLFLG, R, IOUT, IN)
        IF ( IFILLFLG.GT.0 ) THEN
            WRITE (IOUT, 9665)
9665  FORMAT (//' ADDITIONAL DEM FILL FOR SURFACE ROUTING CASCADES ',
+          'WILL BE DONE '//)
            ELSE IF ( IFILLFLG.LE.0 ) THEN
                WRITE (IOUT, 9666)IN
9666  FORMAT (//' ADDITIONAL DEM FILL FOR SURFACE ROUTING CASCADES ',
+          ' WILL NOT BE DONE '//)
            END IF
            ALLOCATE (IPSZBOUND(NCOL,NROW))
            IPSZBOUND=0
            CALL ACTIVECELL
C
C3-----ALLOCATE SPACE FOR SOIL ZONE FLOW.
        IUZM = NROW*NCOL
        ALLOCATE (NUMACTIVE)
        ALLOCATE (DEM2(NCOL,NROW),DIFF(NCOL,NROW))
        ALLOCATE (HASZ(NCOL,NROW),THRSZ(NCOL,NROW))
        ALLOCATE (STRMCELL(NCOL,NROW))
        ALLOCATE (CHEZC(NCOL,NROW), SEEPOLD(NCOL, NROW))
        ALLOCATE (THSSZ(NCOL,NROW),EPSSZ(NCOL,NROW))
        ALLOCATE (FKZSZ(NCOL,NROW),HROFF(NCOL,NROW))
        ALLOCATE (ALPHASZ(NCOL,NROW),THTOSZ(NCOL,NROW))
        ALLOCATE (THTISZ(NCOL,NROW),RUN2STRM(NCOL,NROW))
        ALLOCATE (RSZ(NCOL,NROW), UZF_REJ(NCOL,NROW))
        ALLOCATE (RSZOLD(NCOL,NROW),R2SZ(NCOL,NROW))
        ALLOCATE (RITR(NCOL,NROW), TOT2CELL(NCOL,NROW))
        ALLOCATE (CUM_INF(NCOL,NROW),CUM_INF_P(NCOL,NROW))
        ALLOCATE (DYNCUM_H(NCOL,NROW),PRCHSTRM(NCOL,NROW))
        ALLOCATE (TIME_H(NCOL,NROW),PONDED(NCOL,NROW))
        ALLOCATE (PETSZ(NCOL,NROW),RUNOFFSZ(NCOL,NROW))
        ALLOCATE (TIME_INFILT(NCOL,NROW),TIME_POND(NCOL,NROW))
        ALLOCATE (IORDER(10,IUZM))
        ALLOCATE (IFLWSTRM(NCOL,NROW),IORDERCHK(IUZM))
        ALLOCATE (FLOWOTPSZ(NCOL,NROW))
        ALLOCATE (HA(IUZM,11),PIVOT(IUZM))
        ALLOCATE (CP(3),RP(4),RUNOLD(NCOL,NROW))
        ALLOCATE (RUNIN(NCOL,NROW,4),RUNOUT(NCOL,NROW,4),
+          HITR(NCOL,NROW,NLAY))
        ALLOCATE (QSTORM(NCOL,NROW))

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```

ALLOCATE (ETOUT(NCOL,NROW), PERCHDIS(NCOL,NROW))
ALLOCATE (SUMFLOW(NCOL,NROW), STRMCHD(NCOL,NROW))
ALLOCATE (RUNOLDT(NCOL,NROW))
STRMCELL=0
QSTORM=0.0
SEEPOLD=0.0
STRMCHD=0.0
HROFF=0.0
PETSZ = 0.0
TOT2CELL = 0.0
PONDED = 0.0
CUM_INF = 0.0
CUM_INF_P = 0.0
DYNCUM_H = 0.0
FLOWOTPSZ = 0.0
TIME_H = 0.0
RUNOFFSZ = 0.0
PRCHSTRM = 0.0
TIME_INFILT = 0.0
TIME_POND = 1.0E20
RSZOLD = 0.0
R2SZ = 0.0
RSZ = 0.0
RITR = 0.0
IFLWSTRM = 0
IORDERCHK = 0
IORDER = 0
THTOSZ = 0.0
THTISZ = 0.0
CP = 0.0
RP = 0.0
RUNOLD = 0.0
HA = 0
NUMACTIVE=0
RUNIN = 0.0
RUNOUT = 0.0
ETOUT = 0.0
PERCHDIS = 0.0
RUN2STRM = 0.0
SUMFLOW = 0.0
DEM2=0.0
DIFF=0.0
UZF_REJ=0.0
NUMCHECK=0
ICONFLAG = 0
HITR=HNEW
RUNOLDT = 0.0
CALL U2DREL(HASZ, ANAME(2), NROW, NCOL, 0, IN, IOUT)
CALL U2DREL(THRSZ, ANAME(3), NROW, NCOL, 0, IN, IOUT)
CALL U2DREL(THSSZ, ANAME(4), NROW, NCOL, 0, IN, IOUT)
CALL U2DREL(EPSSZ, ANAME(5), NROW, NCOL, 0, IN, IOUT)
CALL U2DREL(FKZSZ, ANAME(6), NROW, NCOL, 0, IN, IOUT)
CALL U2DREL(ALPHASZ, ANAME(9), NROW, NCOL, 0, IN, IOUT)
CALL U2DREL(THTISZ, ANAME(10), NROW, NCOL, 0, IN, IOUT)
CALL U2DREL(CHEZC, ANAME(12), NROW, NCOL, 0, IN, IOUT)

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C4-----INITIALIZE THETA OLD AND DEM2 FOR FILL
  DO IR = 1, NROW
    DO IC = 1, NCOL
      DEM2(IC,IR)= BOTM(IC,IR,0)
      IF ( IPSZBOUND(IC,IR).GT.0 ) THEN
        THTOSZ(IC,IR) = THTISZ(IC,IR)+1.0E-16
        NUMCHECK = NUMCHECK + 1
      END IF
    END DO
  END DO
C-----SAVE POINTERS FOR GRID AND RETURN
  CALL SGWF2PSZ1PSV(IGRID)
C
C5-----RETURN
  RETURN
END SUBROUTINE GWF2PSZ1AR
C
C-----SUBROUTINE GWF2PSZ1RP
C
  SUBROUTINE GWF2PSZ1RP(IN, KPER, KSTP, IGRID, IUNITSFR)
  USE GLOBAL,          ONLY: NCOL, NROW, IOUT, IBOUND, DELR, DELC
  USE GWFPSZMODULE,   ONLY: PETSZ, RSZ, NUMACTIVE, PERCHFLG,
IFLWSTRM,
  + HROFF
  USE GWFUZFMODULE,   ONLY: VKS
  USE GWFSFRMODULE,   ONLY: NSTRM, ISTRM
  IMPLICIT NONE

C -----
C SPECIFICATIONS:
C -----
C ARGUMENTS
C -----
  INTEGER IN, KPER, IGRID, KSTP, IUNITSFR
C -----
C LOCAL VARIABLES
C -----
  INTEGER NPSZ, IC, IR, L
  CHARACTER*24 ANAME(2)
  DATA ANAME(1)/ ' WATER APPLICATION RATE '/
  DATA ANAME(2)/ '          POTENTIAL ET RATE'/
C -----
C1-----SET POINTERS FOR THE CURRENT GRID.
  CALL SGWF2PSZ1PNT(IGRID)
C
C-----DETERMINE ORDER OF RUNOFF TO CELLS.
  IF (KPER.EQ.1) THEN
    CALL SOILZONE_ORDER()
C
C-----DETERMINE WHAT SOIL ZONE CELLS FLOW TO A STREAM.
  IF ( IUNITSFR.GT.0 ) THEN
    DO L = 1 , NSTRM
      IR = ISTRM(2, L)
      IC = ISTRM(3, L)
      IFLWSTRM(IC, IR) = L
    END DO
  END IF

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```

      END DO
      END IF
    END IF
  C
  C2-----READ IN ARRAY FOR WATER APPLICATION RATE.
    READ (IN, *) NPSZ
    IF ( NPSZ.LT.0 ) THEN
      WRITE (IOUT, *) 'USING WATER APPLICATION RATE FROM PREVIOUS'
      WRITE (IOUT, *) 'STRESS PERIOD.', 'CURRENT PERIOD IS: ', KPER
    ELSE
      CALL U2DREL(RSZ, ANAME(1), NROW, NCOL, 0, IN, IOUT)
    END IF
  C
  C2-----READ IN ARRAY FOR PET RATE.
    READ (IN, *) NPSZ
    IF ( NPSZ.LT.0 ) THEN
      WRITE (IOUT, *) 'USING POTENTIAL ET RATE FROM PREVIOUS STRESS '
      WRITE (IOUT, *) 'PERIOD.', 'CURRENT PERIOD IS: ', KPER
    ELSE
      CALL U2DREL(PETSZ, ANAME(2), NROW, NCOL, 0, IN, IOUT)
    END IF
  C
  C27-----RETURN.
    RETURN
  END SUBROUTINE GWF2PSZ1RP
  C
  C-----SUBROUTINE GWF2PSZ1FM
  C
    SUBROUTINE GWF2PSZ1FM(KKPER, KKSTP, KKITER, IUNITSFR, IUNITLAK,
+      IUNITUZF, IGRID)
    USE GLOBAL,      ONLY: NCOL, NROW, NLAY, IOUT, DELR, DELC,CV,
+      IBOUND, BOTM, ISSFLG, BUFF, RHS, HNEW,
+      HCOF, HOLD
    USE GWFBASMODULE, ONLY: DELT, TOTIM, ICBCFL, TOTIM, PERTIM
    USE GWFPSZMODULE
    USE GWFUFZMODULE, ONLY: FINF, IUZFBND, SEEPOUT, REJ_INF
    USE GWF5FRMODULE, ONLY: NSS, NSTRM, ISTRM, SEG, STRM
    IMPLICIT NONE
  C
  C-----
  C SPECIFICATIONS:
  C-----
  C FUNCTIONS
  C-----
    DOUBLE PRECISION FUNC3, FUNC9, FUNC1, FUNC2, FUNC4, FUNC5, FUNC7
    EXTERNAL FUNC1
    EXTERNAL FUNC2
    EXTERNAL FUNC3
    EXTERNAL FUNC4
    EXTERNAL FUNC5
    EXTERNAL FUNC7
    EXTERNAL FUNC9
  C
  C-----
  C ARGUMENTS
  C-----

```

```

      INTEGER KKPER, KKSTP, IUNITSFR, IUNITLAK, IGRID, IUNITUZF,
KKITER,
      + ICONFLAG
C -----
C LOCAL VARIABLES
C -----
      DOUBLE PRECISION DYNCUM_P, DYNCUM , BETA,
      + ALPHA, FKS, FLOWIN, R, FKI, FKD,
      + TIME, DELTH, SURF, TP, ROLD, DELT_TEMP,
      + CLOSEZERO, CLENGTH, CWIDTH, DELH,
      + R2, FKO, TIME_END, DYNCUM_SAVE, DIFFSEEP,
      + CUM_INF_TEMP, CUM_INF_P_TEMP,
      + TIME_TEMP, CP1, THOCHECK, DYNCUM_H_TEMP, ETACT,
      + TIME_H_TEMP, DEPTHCHECK, FA, FB, FTEMP,
      + TIME_TEMP_A, TIME_TEMP_B, CUM_INF_A, CUM_INF_B, CUM_INF_TEMP_OLD,
      + TIME_TEMP1, F1, F2, TIME_TEMP2, FDERIV, TOL2, ERRORP, SQD
      DOUBLE PRECISION RUNERR, DLR, THI, AA, BB, BIGG, THS, THR, THO,
      + THO_TM1, A, B
      INTEGER I, J, K, II, IC, IR, IK, IBD1, IBD2, IBDPSZ, IRUN,
      + ISS, IFR, IJK, LL, IL, ILL, ICE, IRE
      CHARACTER*16 TEXTPRCH, TEXTET, TEXTINF, TEXTTHO, TEXTRUN
      DATA TEXTRUN/ ' SURF. RUNOFF ' /
      DATA TEXTINF/ ' INFILT. RATE ' /
      DATA TEXTTHO/ ' SURF. SAT. ' /
      DATA TEXTET/ ' SOILZONE ET ' /
      DATA TEXTPRCH/ ' PERCH THCKNESS ' /
C -----
C
C1-----SET POINTERS AND INITIALIZE VALUES FOR THE CURRENT GRID.
      CALL SGWF2PSZ1PNT(IGRID)
      CLOSEZERO = 1.0E-20
      DLR = 1.0E-5
      ISS = ISSFLG(KKPER)
      RUNERR = 0.0
      ICONFLAG = 0
      DO I=1, NSTRM
         STRM(31,I)=0.0
      END DO
      A = 0.0
      B = 0.0
      IF( KKITER.EQ.1 )RITR = 0.0
      LL = 1
      IFR = 0
      IJK = 1
      R = 0.0
      HROFF=0.0
      UZF_REJ=0.0
      FLOWOTPSZ=0.0
      QSTORM=0.0
C
C2-----INDIVIDUAL CELL INFILTRATION, TP, RUNOFF CALCULATION
C
      CONVERGE2: DO WHILE( LL.LE.100 )
         IF ( ICONFLAG.EQ.0) THEN
            DO IR = 1, NROW

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      DO IC = 1, NCOL
        IL=IPSZBOUND(IC, IR)
        !NOTE: IF SZF ALLOWED IN MULTIPLE LAYERS THEN CHANGE
TO(BOUND(IC,IR,IPSZBND(IC, IR))
        !HAVE TO ADJUST SO THAT PERCHED AQUIFERS CAN
        !BE SIMULATED IF THERE ARE ACTIVE CELLS BELOW AN INACTIVE CELL
      IF ( IBOUND(IC,IR,1).NE.0 ) THEN
        DIFFSEEP=0.0
        !INITIALIZE DIFFERENCE BETWEEN SEEPOUT ITERATIONS
        RP(1)=THRSZ(IC,IR)
        RP(2)=THSSZ(IC,IR)
        RP(3)=EPSSZ(IC,IR)
        RP(4)=FKZSZ(IC,IR)
        CP(1)=THRSZ(IC,IR)
        CP(2)=THSSZ(IC,IR)
        CP(3)=HASZ(IC,IR)
        THS=CP(2)
        THR=RP(1)
        BETA = 1.0
        ALPHA = ALPHASZ(IC,IR)
        FKS=RP(4)
        R = RITR(IC,IR) + RSZ(IC,IR)
        ROLD = RSZOLD(IC,IR)
        THI = THTISZ(IC,IR)
        THO = THTOSZ(IC,IR)+1.0E-7
        DELTH = THO - THI
      IF ( R.GE.FKS ) DELTH = THS - THI
        FKI = FUNC3(THI)
        FKO = FUNC3(THO)
        FKD = FKS - FKI
        TIME = TIME_INFILT(IC,IR)
        TIME_END = TIME + DELT
        TP = TIME_POND(IC,IR)
        DYNCUM = CUM_INF(IC,IR) - FKI*TIME
        DYNCUM_SAVE = DYNCUM
        CUM_INF_TEMP = CUM_INF(IC,IR)
        CUM_INF_P_TEMP = CUM_INF_P(IC,IR)
        DYNCUM_H_TEMP = DYNCUM_H(IC,IR)
        TIME_H_TEMP = TIME_H(IC,IR)
C
C3-----CALCUATE MATRIC FLUX POTENTIAL
C
C4-----HIATUS IN STORM.
      IF ( (R-FKS).LT.0.0 ) THEN
        B=THO
        A=THI
        BIGG = FUNC9(A,B) !NOTE: INTEGRATE B&C USING FUNC9
        TP = 1.0E20
        THOCHECK = FUNC7(FKS,R)
        !NOTE: BC WATER CONTENT AT FKS
C
C5-----R<ROLD
C
      IF ( R-ROLD.LT.CLOSEZERO .OR.
        THOCHECK-THO.LT.-CLOSEZERO ) THEN

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```

TIME_H(IC,IR) = TIME
DYNCUM_H(IC,IR) = DYNCUM
TIME_H_TEMP = TIME
DYNCUM_H_TEMP = DYNCUM
THO_TM1 = THO
THO = FUNC5(ALPHA,FKS,THI,THO_TM1,R,TIME,
            TIME_H_TEMP,BETA,DYNCUM_H_TEMP,
            IC,IR)
1
2
ELSE
  THO = THOCHECK
END IF
CUM_INF_TEMP = CUM_INF_TEMP + R * DELT + FKI * DELT
R2 = R
ROLD = R
TIME = TIME + DELT
C6-----DURING STORM EVENT.
ELSE
C7-----IF NOT SATURATED AT BEGINING OF TIMESTEP (ROLD<FKS)
IF( ROLD-FKS.LT.-CLOSEZERO ) THEN
  CUM_INF_TEMP = 0.0
  IF(THO-THS.LT.-CLOSEZERO)THI = THO
  THO = THS
  DELTH = THO - THI

  FKI = FUNC3(THI)
  FKO = FUNC3(THO)
  FKD = FKS - FKI
  TIME = 0.0
  CUM_INF_TEMP = 0.0
  TIME_END = DELT
  TP = 1.0E20
  DYNCUM = 0.0
  DYNCUM_SAVE = 0.0
  END IF
  BIGG = FUNC9(THI,THO)
C8-----CALCULATE PONDING TIME FOR LATEST R.
IF ( TIME.LT.TP.OR.TIME.LT.CLOSEZERO ) THEN
  TP = FUNC4(FKS,ALPHA,R,BIGG,FKI,TIME,DELTH,
            CUM_INF_TEMP,CLOSEZERO)
  IF ( TIME-TP.LT.-CLOSEZERO ) THEN !NOTE: TIME<TP
C9-----INCREMENT TIME BEFORE PONDING.
  IF ( TIME_END.GE.TP ) THEN
    CUM_INF_TEMP = CUM_INF_TEMP + R * TP +
                  FKI * TP
    CUM_INF_P_TEMP = CUM_INF_TEMP
    CUM_INF_P(IC,IR) = CUM_INF_P_TEMP
    TIME = TP
    DYNCUM = CUM_INF_TEMP - FKI * (TP-TIME)
  ELSE !NOTE:TIME END<TP
    CUM_INF_TEMP = CUM_INF_TEMP + R * DELT +
                  FKI * DELT
    DYNCUM = CUM_INF_TEMP - FKI * DELT
    TIME = TIME + DELT
  END IF
END IF
1
2

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```

                END IF
                IF ( TIME-TIME_END.LT.-CLOSEZERO ) THEN
C
C10-----INCREMENT TIME AFTER PONDING.
C
                CUM_INF_TEMP = FUNC2(ALPHA,DELTH,BIGG,FKS,FKD,FKI,
                CUM_INF_TEMP,CUM_INF_P_TEMP,TP,
                TIME_END,DYNCUM,IC,IR,LL, THO, THI)

                TIME = TIME_END
                END IF
                R2 = (DYNCUM-DYNCUM_SAVE)/DELT
                HROFF(IC,IR)=(R-R2)*(DELR(IC)*DELC(IR))
                END IF
                R2SZ(IC,IR) = (R2*DELR(IC)*DELC(IR))
                FLOWOTPSZ(IC,IR)= HROFF(IC,IR)+ SEEPOUT(IC,IR)
                !NOTE: ADD BACK IN REJ_INF ONCE UZF SET UP FOR LAYERS
                IF ( ICONFLAG.EQ.1 .OR. LL.EQ.100 ) THEN
                THTISZ(IC,IR) = THI
                THTOSZ(IC,IR) = THO
                TIME_INFILT(IC,IR) = TIME
                TIME_POND(IC,IR) = TP
                CUM_INF(IC,IR) = CUM_INF_TEMP
                END IF
                END IF
                SEEPOLD(IC,IR)=SEEPOUT(IC,IR)
                END DO
                END DO
C
C11-----END OF INDIVIDUAL CELL COMPUTATION
C
                CALL RUNOFF(IUNITSFR,IFR,IJK,RUNERR,ICONFLAG, KKPER, KKITER,
                LL, ICE, IRE)
                ELSE
                EXIT CONVERGE2
                END IF
                LL = LL + 1
                END DO CONVERGE2
C
C12-----END OF SOIL ZONE CONVERGENCE LOOP DISTRIBUTE SZ FLUXES
C
                CALL ADDFINF(IUNITUZF, KKPER)
                IF ( LL.GE.100) THEN
                WRITE (IOUT, 9002) RUNERR
                WRITE (IOUT,*) ICE,IRE
                9002 FORMAT (1X, 'CONVERGENCE FAILURE IN RUNOFF (ERROR = ',
                E20.10, ') -- SIMULATION ABORTING')
                END IF
                DO IC = 1, NCOL
                DO IR = 1, NROW
                IRUN = IFLWSTRM(IC, IR)
                IF ( IRUN.GT.0 .AND. IUNITSFR.GT.0 ) THEN
                STRM(31, IRUN) = STRM(31, IRUN)+ RUN2STRM(IC,IR)
                END IF
                IF ( ICONFLAG.EQ.1 ) THEN

```

```

          IF ( IBOUND(IC,IR,1).GT.0 ) THEN
            RSZOLD(IC,IR) = RITR(IC,IR) + RSZ(IC,IR)
            RUNOLDT(IC,IR) = RUNOFFSZ(IC,IR)
          END IF
        END IF
      END DO
    END DO
  C
  END SUBROUTINE GWF2PSZ1FM
  C
  C-----SUBROUTINE GWF2PSZ1BD
  C
    SUBROUTINE GWF2PSZ1BD(KKPER, KKSTP, IUNITUZF, IUNITSF, IGRID)
  CR,
    USE GLOBAL,          ONLY: NCOL, NROW, NLAY, IOUT, IBOUND,CV, CC,
+
+           BUFF, DELR, DELC, BOTM, HNEW, RHS
    USE GWF2PSZ1MODULE, ONLY: MSUM, ICBCFL, DELT, PERTIM, TOTIM,
  1
+           VBVL, VBNM
    USE GWF2PSZ1MODULE, ONLY: NSS, NSTRM, ISTRM, SEG, STRM, HSTRM
    USE GWF2PSZ1MODULE, ONLY: FINF, IUZFBND, VKS, REJ_INF
    USE GWF2PSZ1MODULE, ONLY: IPSZCB1, IPSZCB2, R2SZ, IPSZBOUND,
+ FLOWOTPSZ, UZF_REJ, RUNOFFSZ, STRMCELL, STRMCHD, QSTORM,
+ RUN2STRM, IFLWSTRM, HROFF
    USE GWF2PSZ1MODULE, ONLY: A, IA, ICELL
    USE GWF2PSZ1MODULE, ONLY: HKUPW, SN
    IMPLICIT NONE
    CHARACTER*19 TEXT
    DOUBLE PRECISION RIN, ROUT, RATIN, RATOUT, CELBOT, DELH, QQ, HH,
+
+           CHECK, TTOP, BBOT, THICK, CONDIF, DH, CHECK2,
+           SSTRLEN, QCHECK, STAGELEV, DDDHORIZ
    INTEGER IBD1, IBD2, IL, IR, IC, IGRID, KKPER, KKSTP, IUNITUZF, I,
+
+           IUNITSF, IRUN, IJ, L
    REAL ZERO, CLOSEZERO, DUM7
    DATA TEXT / 'INTERFLOW TO SFR' /
  C
  C-----
  C ARGUMENTS
  C-----
    CALL SGWF2PSZ1PNT(IGRID)

  C1-----CLEAR RATIN AND RATOUT ACCUMULATORS, AND SET CELL-BY-CELL
  C1-----BUDGET FLAG.
    CLOSEZERO=1.0E-2
    ZERO=0.
    RATIN=ZERO
    RATOUT=ZERO
    IBD1=0
    IBD2=0
    QQ=0
    QSTORM=0.0
    QCHECK=0.0
    DUM7=0.0 !ALLOW MULT SZ LAY
    IF(IPSZCB1.GT.0 .AND. ICBCFL.NE.0) IBD1=1
    IF(IPSZCB2.GT.0 .AND. ICBCFL.NE.0) IBD2=1
  C

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```

C3-----CLEAR THE BUFFER.
  DO 50 IL=1,NLAY
  DO 50 IR=1,NROW
  DO 50 IC=1,NCOL
  BUFF(IC,IR,IL)=ZERO
50  CONTINUE
C4-----SUBTRACT LEAKAGE TO UZF FROM PERCHED WATER TABLE.
  !NOTE: NEED TO ADD RATOUT FOR SZ DISCHARGE TO LS
  DO IR=1, NROW
  DO IC=1, NCOL
  CHECK=0.0
  IF (KKPER.GT.3) THEN
  CHECK=0.0
  END IF
  IF (IL.GT.0) THEN
  IF (IL+1.LE.NLAY) THEN
  IF (IUZFBND(IC,IR).NE.0) THEN
  FINF(IC,IR)=R2SZ(IC,IR)/(DELC(IR)*DELR(IC))
  IF (FINF(IC,IR).GT.VKS(IC,IR)) THEN
  UZF_REJ(IC,IR)=FINF(IC,IR)-VKS(IC,IR)
  FINF(IC,IR)=VKS(IC,IR)
  REJ_INF(IC,IR)=REJ_INF(IC,IR)+UZF_REJ(IC,IR)
  END IF
  END IF
  END IF
  END IF
  IF (STRMCELL(IC,IR).GT.0) THEN
  L=STRMCELL(IC,IR)
  STAGELEV=HSTRM(L,1)
C5-----CALCULATE QSTORM AND UPDATE RATOUT
  DO IL=1,1 !NOTE IF .GT. 1 ALLOW MULTIPLE SZ LAYER
  IF (STAGELEV.LT.HNEW(IC,IR,IL)) THEN
  IF (IBOUND(IC,IR,IL).GT.0) THEN
  IF (HNEW(IC,IR,IL).GT.BOTM(IC,IR,IL)) THEN
  IF (STAGELEV.GT.BOTM(IC,IR,IL)) THEN
  HH= HNEW(IC,IR,IL)
  TTOP=(BOTM(IC,IR,IL-1))
  BBOT=STAGELEV
  THICK=TTOP-BBOT
  SSTRLEN=2.0*STRM(1,L)
  CONDIF= (HKUPW(IC,IR,IL)*SSTRLEN)/
  +          ((DELC(IR)+DELR(IC))/4)
  DH=DDDORIZ(HH,TTOP,BBOT,IL) !DSN/DH
  IJ=ICELL(IC,IR,IL)
  QCHECK=(CONDIF*SN(IJ)*THICK*(STAGELEV-HH))
  ELSE
  HH= HNEW(IC,IR,IL)
  TTOP=(BOTM(IC,IR,IL-1))
  BBOT=(BOTM(IC,IR,IL))
  THICK=TTOP-BBOT
  SSTRLEN=2*STRM(1,L)
  CONDIF= (HKUPW(IC,IR,IL)*SSTRLEN)/
  +          ((DELC(IR)+DELR(IC))/4)
  DH=DDDORIZ(HH,TTOP,BBOT,IL) !DSN/DH
  IJ=ICELL(IC,IR,IL)

```

```

        QCHECK=(CONDIF*SN(IJ)*THICK*(BBOT-HH))
        DUM7=DUM7+QCHECK
        END IF
        !NOTE: PREVENTS BACKFLOW INTO SZ; FLOW INTO SZ HANDLED
        BY GFE AND CONDUCTANCE THROUGH MF FLOW PKG
        IF (QSTORM(IC,IR).LT.0.0) QSTORM(IC,IR)=0.0
        END IF
        END IF
        END IF
        END DO
        QSTORM(IC,IR)=QSTORM(IC,IR)+DUM7
        QQ=QSTORM(IC,IR)
        RATOUT=RATOUT-QQ
        END IF
        END DO
        END DO
C6-----SET VALUE FOR Q AND UPDATE BUFFER
        IF (IBD1.GT.0) THEN
            DO IR=1, NROW
                DO IC=1, NCOL
                    !NOTE: NEED TO RE-EVAL IF SZF NOT L1
                    IF(IPSZBOUND(IC, IR).EQ.0) THEN
                        IL=1
                    ELSE
                        IL=IPSZBOUND(IC,IR)
                    END IF
                    BUFF(IC,IR,IL)=BUFF(IC,IR,IL)+ R2SZ(IC,IR)
                END DO
            END DO
            CALL UBUDSV(KKSTP, KKPER, TEXT, IPSZCB1, BUFF,
+                               NCOL, NROW, NLAY, IOUT)
        END IF
        IF (IBD2.GT.0) THEN
            DO IR=1, NROW
                DO IC=1, NCOL
                    IF(IPSZBOUND(IC, IR).EQ.0) THEN
                        IL=1
                    ELSE
                        IL=IPSZBOUND(IC,IR)
                    END IF
                    BUFF(IC,IR,IL)=BUFF(IC,IR,IL)+ R2SZ(IC,IR)
                END DO
            END DO
            CALL UBDSV3(KKSTP, KKPER, TEXT, IPSZCB2,
+                               BUFF, IPSZBOUND, 1, NCOL, NROW,
+                               NLAY, IOUT, DELT, PERTIM, TOTIM,
+                               IBOUND)
        END IF
C7-----MOVE RATES, VOLUMES & LABELS INTO ARRAYS FOR PRINTING.
!THIS SETS UP BUDGET PRINTING TEXT=BUDGET ITEM
        VBVL(4,MSUM)=RATOUT
        VBVL(2,MSUM)=VBVL(2,MSUM)+RATOUT*DELT
        VBNM(MSUM)=TEXT
C
C8-----INCREMENT BUDGET TERM COUNTER(MSUM).

```

```

MSUM=MSUM+1
C
C9-----WRITE STATEMENT TO PRINT OUT INTERFLOW AND SEEPAGE TO UZF
  WRITE (81,*) KKPER
  WRITE (81,*)
  DO IR=1,NROW
    DO IC=1,NCOL
      IF(QSTORM(IC,IR).GT.0.0) THEN
        WRITE(81,*)KKPER, QSTORM(3,1)
      END IF
    END DO
  END DO
  WRITE (82,*) KKPER
  WRITE (82,*)
  DO IR=1,NROW
    DO IC=1,NCOL
      IF(FINF(IC,IR).GT.0.0) THEN
        WRITE(82,999) FINF(IC,IR)
      END IF
    END DO
  END DO
999  FORMAT(1000E10.5)
1111 FORMAT(81F8.2)
  RETURN
  END SUBROUTINE GWF2PSZ1BD
C
C-----SUBROUTINE RUNOFF
C
  SUBROUTINE RUNOFF(IUNITSFR,IFR,IJK,FERROR,ICONFLAG,KKPER,KKITER,
+ LL, ICE,IRE)!WH ADD KKITER,LL FOR DEBUG
  USE GLOBAL, ONLY: NCOL, NROW, NLAY, IOUT, DELR, DELC,
+ IBOUND, BOTM
  USE GWFBASMODULE, ONLY: DELT, TOTIM, ICBCFL, TOTIM, PERTIM
  USE GWFPSZMODULE, ONLY: NUMACTIVE, IORDER, DEM2, FLOWOTPSZ,
+ CHEZC, RUNIN, RUNOUT, RITR, RUNOFFSZ,
+ IFLWSTRM, RUN2STRM, RUNOLD
  USE GWFSFRMODULE, ONLY: NSS, NSTRM, ISTRM, STRM
C
-----
  IMPLICIT NONE
C
-----
C
  ARGUMENTS
C
-----
  DOUBLE PRECISION CLOSEZERO, FERROR, QA, QB, QC, QD, SLOPE, WIDTH,
+ C, DELX, TOL, FPART,DELZ, DELY, FTEST
  INTEGER IFR, NUMLEFT,NUMITR, II, IC, IR, TOTRUN, ICM, ICP, IRM,
+ IRP, IFLIC, IFLIR,ICONFLAG, ICE, IRE, IUNITSFR, IJK,
+ KKPER,KKITER, IRUN, RUNCON,LL,DUM
  REAL AREA, SLOPEMIN, ELEVLOW
C
-----
  IFR = 0
  NUMLEFT = NCOL*NROW
  NUMITR = 1
  CLOSEZERO = 1.0E-20
  TOL = 5.0E-3
  SLOPEMIN = 1.0E-6

```

```

RUNCON=0
FTEST=0.0
C1-----LOOP THROUGH CALCULATION ORDER.
FERROR=0.0
RITR=0.0 !NOTE: RESET EVERY FORMULATE
DO II = 1, NUMACTIVE
  IC = IORDER(1,II) !NOTE:THESE ARE SET IN SZ ORDER SBRT
  IR = IORDER(2,II)
  AREA = DELR(IC)*DELC(IR)
  TOTRUN = IORDER(7,II)+IORDER(8,II)+IORDER(9,II)+IORDER(10,II)
  ICM = 0
  ICP = 0
  IRM = 0
  IRP = 0
  ICE = IC
  IRE = IR
  IFLIC = IC
  IFLIR = IR
  ELEVLOW = DEM2(IC, IR)
  RUNOFFSZ(IC,IR)=0.0
  IF ( TOTRUN.GT.0 ) THEN
    FPART = FLOWOTPSZ(IC,IR)/TOTRUN
    IF ( FPART.GT.CLOSEZERO ) THEN
      IF( IC.GT.1 .AND. IORDER(7,II).GT.0 )THEN
        WIDTH = DELC(IR)
        DELX = (DELR(IC)+DELR(IC-1))/2.0
        DELZ = (DEM2(IC,IR) - DEM2(IC-1,IR))
        SLOPE = DELZ/DELX
        IF ( SLOPE.LT.SLOPEMIN ) SLOPE = SLOPEMIN
        C = (CHEZC(IC,IR)+CHEZC(IC-1,IR))/2.0
        QA = RUNIN(IC-1,IR,1)
        QB = RUNOUT(IC-1,IR,1)
        QC = IORDER(7,II)*FPART
        CALL ROUTE_CHAN2(QA, QB, QC, QD, DELT, SLOPE, WIDTH, C,
          + DELX, IC, IR)
        RITR(IC-1,IR) = RITR(IC-1,IR) + QD/AREA
        RUNOFFSZ(IC-1,IR) = RUNOFFSZ(IC-1,IR) + QD
        IF ( ICONFLAG.EQ.1 ) RUNIN(IC-1,IR,1) =
IORDER(7,II)*FPART
        IF ( ICONFLAG.EQ.1 ) RUNOUT(IC-1,IR,1) = QD
      END IF
      IF( IC.LT.NCOL .AND. IORDER(8,II).GT.0 ) THEN
        WIDTH = DELC(IR)
        DELX = (DELR(IC)+DELR(IC+1))/2.0
        DELZ = (DEM2(IC,IR) - DEM2(IC+1,IR))
        SLOPE = DELZ/DELX
        IF ( SLOPE.LT.SLOPEMIN ) SLOPE = SLOPEMIN
        C = (CHEZC(IC,IR)+CHEZC(IC+1,IR))/2.0
        QA = RUNIN(IC+1,IR,2)
        QB = RUNOUT(IC+1,IR,2)
        QC = IORDER(8,II)*FPART
        CALL ROUTE_CHAN2(QA, QB, QC, QD, DELT, SLOPE, WIDTH, C,
          + DELX)
        RITR(IC+1,IR) = RITR(IC+1,IR) + QD/AREA
        RUNOFFSZ(IC+1,IR) = RUNOFFSZ(IC+1,IR) + QD
      END IF
    END IF
  END IF
END DO

```

```

      IF ( ICONFLAG.EQ.1 ) RUNIN(IC+1,IR,2) =
IORDER(8,II)*FPART
      IF ( ICONFLAG.EQ.1 ) RUNOUT(IC+1,IR,2) = QD
      END IF
      IF( IR.GT.1 .AND. IORDER(9,II).GT.0 ) THEN
        WIDTH = DELR(IC)
        DELY = (DELIC(IC)+DELIC(IC-1))/2.0
        DELZ = (DEM2(IC,IR) - DEM2(IC,IR-1))
        SLOPE = DELZ/DELY
        IF ( SLOPE.LT.SLOPEMIN ) SLOPE = SLOPEMIN
        C = (CHEZC(IC,IR)+CHEZC(IC,IR-1))/2.0
        QA = RUNIN(IC,IR-1,3)
        QB = RUNOUT(IC,IR-1,3)
        QC = IORDER(9,II)*FPART
        CALL ROUTE_CHAN2(QA, QB, QC, QD, DELT, SLOPE, WIDTH, C,
+                               DELY)
        RITR(IC,IR-1) = RITR(IC,IR-1) + QD/AREA
        RUNOFFSZ(IC,IR-1) = RUNOFFSZ(IC,IR-1) + QD
        IF ( ICONFLAG.EQ.1 ) RUNIN(IC,IR-1,3) =
IORDER(9,II)*FPART
        IF ( ICONFLAG.EQ.1 ) RUNOUT(IC,IR-1,3) = QD
        END IF
        IF( IR.LT.NROW .AND. IORDER(10,II).GT.0 ) THEN
          WIDTH = DELR(IC)
          DELY = (DELIC(IC)+DELIC(IC+1))/2.0
          DELZ = (DEM2(IC,IR) - DEM2(IC,IR+1))
          SLOPE = DELZ/DELY
          IF ( SLOPE.LT.SLOPEMIN ) SLOPE = SLOPEMIN
          C = (CHEZC(IC,IR)+CHEZC(IC,IR+1))/2.0
          QA = RUNIN(IC,IR+1,4)
          QB = RUNOUT(IC,IR+1,4)
          QC = IORDER(10,II)*FPART
          CALL ROUTE_CHAN2(QA, QB, QC, QD, DELT, SLOPE, WIDTH, C,
+                               DELY)
          RITR(IC,IR+1) = RITR(IC,IR+1) + QD/AREA
          RUNOFFSZ(IC,IR+1) = RUNOFFSZ(IC,IR+1) + QD
          IF ( ICONFLAG.EQ.1 )
+
          RUNIN(IC,IR+1,4) = IORDER(10,II)*FPART
          IF ( ICONFLAG.EQ.1 ) RUNOUT(IC,IR+1,4) = QD
          END IF
        END IF
      ELSE
        RUNOFFSZ(IC,IR) = FLOWOTPSZ(IC,IR)
      END IF
    END DO
  DO IR = 1, NROW
  DO IC = 1, NCOL
    IF ( IBOUND(IC,IR,1).NE.0 ) THEN !NOTE:CHANGED TO NE
      IRUN = IFLWSTRM(IC, IR) !IRUN=STRMCELL OF DISCHARGE
      IF ( IRUN.GT.0 .AND. IUNITSFR.GT.0 ) THEN
        RUN2STRM(IC,IR) = RUNOFFSZ(IC,IR)
      END IF
      FTEST = ABS(RUNOFFSZ(IC,IR) -
RUNOLD(IC,IR))!NOTE:RUNOLD=LAST ITER RUNOFF
      IF ( FTEST.GT.FERROR+1.0E-7 ) THEN

```

```

                !NOTE:DETERMINE WHICH CELL IS MOST PROBLEMATIC
                FERROR = FTEST
                ICE = IC
                IRE = IR
            END IF
        END IF
        RUNOLD(IC,IR) = RUNOFFSZ(IC,IR)
    END DO
END DO
IF ( FERROR.LT.TOL ) ICONFLAG = 1
IFR = 1
RETURN
END SUBROUTINE RUNOFF

C
C-----SUBROUTINE ROUTE_CHAN2
C
    SUBROUTINE ROUTE_CHAN2(QA, QB, QC, QD, DELTINC, SLOPE, WIDTH1, C,
+                               DEL, IC, IR)
C*****
*
C    IMPLICIT FINITE-DIFFERENCE SCHEME TO ROUTE OVERLAND FLOW
C    VERSION 2.6: DECEMBER 6, 2005
C*****
*
    USE GLOBAL,          ONLY: IOUT
    USE GWFPSZMODULE     !PERHAPS CONSIDER MAKLING AN ONLY CLAUSE

C
C-----
C
C    SPECIFICATIONS:
C-----
C
C    ARGUMENTS
C-----
C
C-----
C
C    IMPLICIT NONE
C    DOUBLE PRECISION SLOPE, WIDTH1, QA, QB, QC, QD, C, DEL, WIDTH
C    REAL DELTINC
C    INTRINSIC ABS, DABS

C-----
C
C    LOCAL VARIABLES
C-----
C
C
C    INTEGER I, MAXITER, IC, IR
C    REAL W_1, WEIGHT
C    DOUBLE PRECISION QDERIV, DQ, DELQ, TOL, AB, AC, AA, AD1,
CLOSEZERO
C    DOUBLE PRECISION DEPTH, QD2, F11, F12, F1, F2, AD2, CONST

C-----
C
C    !NOTE: QA QB RUNIN RUNOUT PREV ITER
C    MAXITER = 100
C    WEIGHT = 0.9
C    CLOSEZERO = 1.0D-20
C    W_1 = 1.0 - WEIGHT
C    DQ = 1.0E-12

```

```

TOL = 1.0E-6
I = 1
WIDTH = 0.1*WIDTH1
C
C2-----MAKE AN INITIAL GUESS AT QD.
QD = (QC+QB)/2.0
C
C3-----INITIALIZE CONSTANTS.
CONST = C*WIDTH*SLOPE**0.5
DELQ = 0.0D0
IF ( CONST.GT.CLOSEZERO ) THEN
  DEPTH = (QA/CONST)**(2.0/3.0)
  AA = DEPTH*WIDTH
  DEPTH = (QB/CONST)**(2.0/3.0)
  AB = DEPTH*WIDTH
  DEPTH = (QC/CONST)**(2.0/3.0)
  AC = DEPTH*WIDTH
ELSE
  AA = WIDTH
  AB = WIDTH
  AC = WIDTH
END IF
C
C4-----CALCULATE FLOW IN CHANNELS--MAXIMUM ITERATIONS IS 50.
CONVERGE: DO WHILE( I.LT.MAXITER )
C
C5-----LOOP THROUGH UNTIL DELQ LESS THAN TOLERANCE.
IF ( DABS(DELQ).GT.TOL .OR. I.EQ.1 ) THEN
  QD2 = QD + DQ
C
C6-----CALCULATE VARIABLES AD1 AND AD2.
IF ( CONST.GT.CLOSEZERO ) THEN
  DEPTH = (QD/CONST)**(2.0/3.0)
  AD1 = DEPTH*WIDTH
  DEPTH = (QD2/CONST)**(2.0/3.0)
  AD2 = DEPTH*WIDTH
ELSE
  AD1 = WIDTH
  AD2 = WIDTH
END IF
C
C7-----CALCULATE FLOW.
F11 = (WEIGHT*((QD)-QC)+(W_1)*(QB-QA))/DEL
F12 = (AD1-AB)/(DELTINC)
F1 = F11+F12
F11 = (WEIGHT*((QD+DQ)-QC)+(W_1)*(QB-QA))/DEL
F12 = (AD2-AB)/(DELTINC)
F2 = F11+F12
QDERIV = (F2-F1)/DQ
DELQ = -F1/QDERIV
IF ( (QD + DELQ).LT.0.0 ) THEN
  IF ( QD-TOL.GT.CLOSEZERO ) THEN
    DELQ = -QD
    QD = 0.0D0
  ELSE

```

```

        QD = 0.0D0
        DELQ = 2.0*TOL
        END IF
    ELSE
        QD = QD + DELQ
    END IF
C
C8-----EXIT LOOP IF DELQ LESS THAN TOLERANCE.
    ELSE
        EXIT CONVERGE
    END IF
    I = I + 1
END DO CONVERGE
C
    IF ( I.GE.MAXITER ) WRITE(IOUT,*)
+           'NON-CONVERGENCE IN OVERLAND FLOW ROUTING',I,QD,
+           DELQ
C
C9-----RETURN.
    RETURN
END SUBROUTINE ROUTE_CHAN2

DOUBLE PRECISION FUNCTION FUNC1(BIGG,DELTH,ALPHA,FKS,DYNCUM)
IMPLICIT NONE
C1-----CALCULATE FC
DOUBLE PRECISION F1, DYNCUM, ALPHA, FKS, DELTH, BIGG
INTRINSIC EXP
F1=DYNCUM*ALPHA/(DELTH*BIGG)
FUNC1=FKS*(1.0+ALPHA/(EXP(F1)-1.0))
END FUNCTION FUNC1
C
DOUBLE PRECISION FUNCTION FUNC2(ALPHA,DELTH,BIGG,FKS,FKD,FKI,
+           CUM_INF_TEMP,CUM_INF_P,TP,
+           TIME_TEMP,DYNCUM,IC, IR,LL, THO, THI)
USE GLOBAL, ONLY: IOUT
USE GWFBASMODULE, ONLY: DELT
IMPLICIT NONE
DOUBLE PRECISION BIGG, F1,F2, CHECK, DCUM,CLOSEZERO,TOL,
+ CUMINC, F3, DELTH, FKS, FKD, F4,ALPHA, GAMMA,DYNCUM_P,
+ CUM_INF_P,FKI, TP, CUM_INF_A, CUM_INF_TEMP, CUM_INFB,
+ TIME_TEMP, DYNCUM, D, CTOP,C, FNUM, FA, FB,F11, THO, THI
INTEGER IC, IR, LL, I
INTRINSIC LOG, ABS
C
C1-----CALCULATE CUMULATIVE INFILTRATION AFTER PONDING
CHECK=0.0 !WH ADD
DCUM = 1.0E-8
CLOSEZERO = 1.0E-20
TOL = 1.0E-5
I = 1
CUMINC = 0.0
F3=(DELTH*BIGG*FKS)/FKD !NOTE: EQN 22 SMITH CORDINI
F4=-FKS*(1.0-ALPHA)
GAMMA=ALPHA*FKS/FKD !SP GAMMA SHAPE FACTOR
DYNCUM_P=CUM_INF_P-FKI*TP

```

```

!NOTE: BOUNDING SOLN BETWEEN A&B FOR BISECTION METHOD
CUM_INFA = 1.01*DYN_CUM_P + FKI*TIME_TEMP
CUM_INF_B = CUM_INF_TEMP + 5.0*FKS*DELT + FKI*DELT
CUM_INF_TEMP = (CUM_INFA + CUM_INF_B)/2.0
C2-----CALCULATE FA AND FB
!FA CALCULATION
IF (CUM_INFA - (FKI*TIME_TEMP).GT.CLOSEZERO) THEN
  DYN_CUM=CUM_INFA - FKI*TIME_TEMP
ELSE
  DYN_CUM=0.0
  CUM_INF_TEMP = (CUM_INFA + CUM_INF_B)/2.0
END IF
F1=ALPHA*DYN_CUM/(DELTH*BIGG)
F2=ALPHA*DYN_CUM_P/(DELTH*BIGG)
D=CUM_INFA-CUM_INF_P !NOTE: I-IP
CTOP = CLOSEZERO
IF ( F1.LT.70. .AND. F1.GT.CLOSEZERO )
!NOTE: ANYTHING HIGHER THAN 70 TOO BIG
+   CTOP = (EXP(F1)-1.0+GAMMA)
C=(CTOP)/(EXP(F2)-1.0+GAMMA)
IF ( C.LT.CLOSEZERO ) C = CLOSEZERO
FNUM=D-F3*LOG(C)
FA=F4*((TIME_TEMP)-TP)+(FNUM) !NOTE: FA CAN BE NEG
!FB CALCULATION
IF ((CUM_INF_B - FKI*TIME_TEMP).GT.CLOSEZERO) THEN
  DYN_CUM=CUM_INF_B - FKI*TIME_TEMP
ELSE
  DYN_CUM= 0.0
  CUM_INF_B=FKI*TIME_TEMP
END IF
F1=ALPHA*DYN_CUM/(DELTH*BIGG)
F2=ALPHA*DYN_CUM_P/(DELTH*BIGG)
D=CUM_INF_B-CUM_INF_P
CTOP = 2.51544E30
IF ( F1.LT.70. .AND. F1.GT.CLOSEZERO )
+   CTOP = (EXP(F1)-1.0+GAMMA)
C=(CTOP)/(EXP(F2)-1.0+GAMMA)
IF ( C.LT.CLOSEZERO ) C = CLOSEZERO
FNUM=D-F3*LOG(C) !NOTE: EQN 22 SMITH CORRADINI
FB=F4*((TIME_TEMP)-TP)+(FNUM)
C3-----INITIAL F11 CALCULATION-CONVERGENCE CRITERIA
IF ((CUM_INF_TEMP - FKI*TIME_TEMP).GT.CLOSEZERO) THEN
  DYN_CUM=CUM_INF_TEMP - FKI*TIME_TEMP
ELSE
  DYN_CUM=0.0
  CUM_INF_TEMP=FKI*TIME_TEMP
END IF
F1=ALPHA*DYN_CUM/(DELTH*BIGG)
F2=ALPHA*DYN_CUM_P/(DELTH*BIGG)
D=CUM_INF_TEMP-CUM_INF_P
CTOP = 2.51544E30
IF ( F1.LT.70. .AND. F1.GT.CLOSEZERO ) !NOTE: EXP(70) TOO BIG
+   CTOP = (EXP(F1)-1.0+GAMMA)
  C=(CTOP)/(EXP(F2)-1.0+GAMMA)
  IF ( C.LT.CLOSEZERO ) C = CLOSEZERO

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```

FNUM=D-F3*LOG(C)
F11=F4*((TIME_TEMP)-TP)+(FNUM) !NOTE: FUNC EVAL WITH INTIAL GUESS
C4-----CHECK RESIDUAL AND ITERATE BISECTION AS NEEDED.
CONVERGE: DO WHILE( I.LT.50 )
  IF ( ABS(F11).GT.TOL ) THEN !NOTE: CONVERGENCE CHECK
    IF ( F11*FB.GT.0.0 ) THEN
      CUM_INF_B = CUM_INF_TEMP
      FB=F11
    ELSE
      CUM_INF_A = CUM_INF_TEMP
      FA=F11
    END IF
    !NOTE: BISECT VALUE
    CUM_INF_TEMP=(CUM_INF_B+CUM_INF_A)/2.0
    !NOTE: RECALCULATE F11 FOR BISECTED VALUE.
    IF((CUM_INF_TEMP - FKI*TIME_TEMP).GT.CLOSEZERO) THEN
      DYNCUM=CUM_INF_TEMP - FKI*TIME_TEMP
    ELSE
      DYNCUM=0.0
      CUM_INF_TEMP=FKI*TIME_TEMP
    END IF
    F1=ALPHA*DYNCUM/(DELTH*BIGG)
    F2=ALPHA*DYNCUM_P/(DELTH*BIGG)
    D=CUM_INF_TEMP-CUM_INF_P
    CTOP = 2.51544E30
    IF ( F1.LT.70. .AND. F1.GT.CLOSEZERO )
      !NOTE:SETS CTOP IF.LT.70
      CTOP = (EXP(F1)-1.0+GAMMA)
      C=(CTOP)/(EXP(F2)-1.0+GAMMA)
      IF ( C.LT.CLOSEZERO ) C = CLOSEZERO
      FNUM=D-F3*LOG(C)
      F11=F4*((TIME_TEMP)-TP)+(FNUM) !NOTE: FUNCTION EVALUATION
    ELSE !NOTE: IF F11.LE.TOL
      DYNCUM=CUM_INF_TEMP - FKI*(TIME_TEMP)
      EXIT CONVERGE
    END IF
    I = I + 1
  END DO CONVERGE
  IF ( I.GE.100 ) THEN
    WRITE(IOUT,*) 'NON-CONVERGENCE FOR TIME AFTER PONDING ',
    'ERROR = ',LL,IC,IR,F11
  END IF
  FUNC2 = CUM_INF_TEMP
END FUNCTION FUNC2
C
DOUBLE PRECISION FUNCTION FUNC3(DUM1)
USE GWFPSZMODULE, ONLY: RP
IMPLICIT NONE
DOUBLE PRECISION:: DUM1
C1-----USE BROOK-COREY TO CALCULATE UNSATURATED K.
IF(DUM1.LT.RP(2)) THEN
  FUNC3=RP(4)*(((DUM1-RP(1))/(RP(2)-RP(1)))**RP(3))
ELSE
  FUNC3=RP(4)
END IF

```

```

END FUNCTION FUNC3

DOUBLE PRECISION FUNCTION FUNC4(FKS,ALPHA,R,BIGG,FKI,TIME,DELTH,
1      CUM_INF_TEMP,CLOSEZERO)
IMPLICIT NONE
DOUBLE PRECISION BIGG, R, FKS, ALPHA, FK1, TIME, DELTH,
+ CUM_INF_TEMP, CLOSEZERO, F1, F2, F3
INTRINSIC LOG, ABS
C1-----CALCULATE TP FOR GIVEN R
IF ( ABS(R-FKS).GT.CLOSEZERO ) THEN
    F1=R*TIME-CUM_INF_TEMP
    F2=BIGG*DELTH/ALPHA
    F3=LOG((R-FKS+ALPHA*FKS)/(R-FKS))
    FUNC4=(F2*F3+F1)/(R-FKI)
ELSE
    FUNC4 = 1.0E20
END IF
END FUNCTION FUNC4

C
DOUBLE PRECISION FUNCTION FUNC5(ALPHA,FKS,THI,THO_TM1,R,TIME,
1      TIME_H,BETA,DYNCUM_H,IC,IR)
USE GLOBAL,      ONLY: IOUT
USE GWFPSZMODULE, ONLY: CP
USE GWFBASMODULE, ONLY: DELT
IMPLICIT NONE
EXTERNAL FUNC3, FUNC9
DOUBLE PRECISION A, B, BIGG, FUNC3, THO, THI, THO_TM1,THS,
+      K1,THOK3,THOK4,THOK5,THOK6, CLOSEZERO,
+      ALPHA,FKS,TIME,TIME_H,H,TE,FCT,
+      K2,K3,K4,YIP1,X,EPS,HNEW,K6,K5, FUNC9,
+      HMIN,HMAX,BETA, DELTHO, FK1, DYNCUM_Z, DYNCUM_H,
+      R, FKO, FUNC, F1, F2, F11
REAL P
INTEGER L,I,MAXITER,KK, IC, IR
C1-----CALCULATE TIME AFTER PONDING
THS=CP(2)
P = 2.0
EPS = 1.0E-6
I = 1
MAXITER = 100
BETA = 0.90 !SMITH PARLANGE BETA
DELTHO = 0.0
CLOSEZERO = 1.0E-16
THO = THO_TM1
IF ( ABS(THO-THI).LT.CLOSEZERO )THO = THI-5.0*EPS
FKI = FUNC3(THI)
X = DELT/3.0
FCT = 0.9
H = DELT
HMAX = DELT
HMIN = DELT/1000.
C2-----BEGIN RUNGE-KUTTA
!K1
CONVERGE: DO WHILE ( X.LT.DELT )
    IF ( I.GE.MAXITER ) THEN

```

```

EXIT CONVERGE
ELSE
  I = I + 1
  IF ( DELT-X.LT.H )H = DELT-X
  DYNCUM_Z = DYNCUM_H + (R-FKI)*(TIME-TIME_H)
  IF ( DYNCUM_Z.LT.EPS ) DYNCUM_Z = EPS
  DELTHO = THO-THI
  B=THO
  FKO=FUNC3(B)
  BIGG = FUNC9(THI,B)
  IF (BIGG.LT.CLOSEZERO)BIGG = CLOSEZERO
  F1 = DELTHO/DYNCUM_Z
  F2 = BETA*P*FKS*DELTHO*BIGG/DYNCUM_Z
  F11 = F1*(R-FKI-FKO-F2)
  K1 = H*F11
  !K2
  DYNCUM_Z = DYNCUM_H + (R-FKI)*(TIME+H/4.0-TIME_H)
  DELTHO = THO+K1/4.0-THI
  B=THO+K1/4.0
  FKO=FUNC3(B)
  BIGG = FUNC9(THI,B)
  IF (BIGG.LT.CLOSEZERO)BIGG = CLOSEZERO
  F1 = DELTHO/DYNCUM_Z
  F2 = BETA*P*FKS*DELTHO*BIGG/DYNCUM_Z
  F11 = F1*(R-FKI-FKO-F2)
  K2 = H*F11
  !K3
  THOK3=THO+(3.0/32.0)*K1+(9.0/32.0)*K2
  DYNCUM_Z = DYNCUM_H + (R-FKI)*(TIME+(3.0/8.0)*H-TIME_H)
  DELTHO = THOK3-THI
  B=THOK3
  FKO=FUNC3(B)
  BIGG = FUNC9(THI,B)
  IF (BIGG.LT.CLOSEZERO)BIGG = CLOSEZERO
  F1 = DELTHO/DYNCUM_Z
  F2 = BETA*P*FKS*DELTHO*BIGG/DYNCUM_Z
  F11 = F1*(R-FKI-FKO-F2)
  K3 = H*F11
  !K4
  THOK4 = THO+(1932./2197.)*K1-
(7200./2197.)*K2+(7296./2197.)*K3
  DYNCUM_Z = DYNCUM_H + (R-FKI)*(TIME+(12./13.)*H-TIME_H)
  DELTHO = THOK4-THI
  B=THOK4
  FKO=FUNC3(B)
  BIGG = FUNC9(THI,B)
  IF (BIGG.LT.CLOSEZERO)BIGG = CLOSEZERO
  F1 = DELTHO/DYNCUM_Z
  F2 = BETA*P*FKS*DELTHO*BIGG/DYNCUM_Z
  F11 = F1*(R-FKI-FKO-F2)
  K4 = H*F11
  !K5
  THOK5 = THO+(439./216.)*K1-8.0*K2+(3680./513.)*K3-
(845./4104.)*K4
  DYNCUM_Z = DYNCUM_H + (R-FKI)*(TIME+H-TIME_H)

```

```

DELTHO = THOK5-THI
B=THOK5
FKO=FUNC3(B)
BIGG = FUNC9(THI,B)
IF (BIGG.LT.CLOSEZERO)BIGG = CLOSEZERO
F1 = DELTHO/DYNCUM_Z
F2 = BETA*P*FKS*DELTHO*BIGG/DYNCUM_Z
F11 = F1*(R-FKI-FKO-F2)
K5 = H*F11
!K6
THOK6 = THO-(8./27.)*K1+2.0*K2-(3544./2565.)*K3+
      (1859./4104.)*K4-(11./40.)*K5
+
DYNCUM_Z = DYNCUM_H + (R-FKI)*(TIME+0.5*H-TIME_H)
DELTHO = THOK6-THI
B=THOK6
FKO=FUNC3(B)
BIGG = FUNC9(THI,B)
IF (BIGG.LT.CLOSEZERO)BIGG = CLOSEZERO
F1 = DELTHO/DYNCUM_Z
F2 = BETA*P*FKS*DELTHO*BIGG/DYNCUM_Z
F11 = F1*(R-FKI-FKO-F2)
K6 = H*F11
TE = ABS((1./360.)*K1-(128./4275.)*K3-(2197./75240.)*K4
      +(1./50.)*K5+(2./55.)*K6)
+
HNEW = FCT*H*(H*EPS/TE)**0.25
IF ( TE.LE.H*EPS )THEN
+
  THO = THO_TM1 +(16./135.)*K1+(6656./12825.)*K3+
      (28561./56430.)*K4-(9.0/50.0)*K5+(2.0/55.0)*K6
  THO_TM1 = THO
  H = HNEW
  IF ( H.LT.HMIN ) H = HMIN
  X = X + H
END IF
IF ( TE.GT.H*EPS )THEN
  H = HNEW
  IF ( H.GT.HMAX ) H = HMAX
END IF
END IF
END DO CONVERGE
IF (I.GE.200) WRITE(IOUT,*) "FAILED IN FUNC5", IC, IR,X,DELT
DYNCUM_Z = DYNCUM_H + (R-FKI)*(TIME-TIME_H)
FUNC5 = THO
END FUNCTION FUNC5
C
DOUBLE PRECISION FUNCTION FUNC7(FKS,R)
USE GWFPSZMODULE, ONLY: RP
IMPLICIT NONE
DOUBLE PRECISION R, FKS
C1-----CALCULATE THETA WHEN R>ROLD AND R<FKS.
FUNC7 = (RP(2)-RP(1))*(R/FKS)**(1.0/RP(3)) + RP(1)
END FUNCTION FUNC7
C
DOUBLE PRECISION FUNCTION FUNC9(A,B)
USE GWFPSZMODULE, ONLY: RP
IMPLICIT NONE

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      DOUBLE PRECISION:: A, B, TERM1, TERM2, TERM_DIFF
C1-----RETURNS INTEGRAL OF BROOKS COREY RELATIVE PERMEABILITY
      EQUATION(SMITH PARLANGE G)
      TERM1=(((B-RP(1))*(RP(1)-B)/(RP(1)-
RP(2)))*(RP(3))/(1+RP(3)))
      TERM2=(((A-RP(1))*(RP(1)-A)/(RP(1)-
RP(2)))*(RP(3))/(1+RP(3)))
      TERM_DIFF=TERM1-TERM2
      FUNC9=RP(4)*TERM_DIFF
      END FUNCTION FUNC9
C
C-----SUBROUTINE SOILZONE_ORDER
C
      SUBROUTINE SOILZONE_ORDER
      USE GWFPSZMODULE, ONLY: IORDER, IORDERCHK, NUMACTIVE, DEM2, DIFF,
+      STRMCELL, IFILLFLG
      USE GLOBAL,      ONLY: NCOL, NROW, NLAY, IBOUND, BOTM, IOUT
      IMPLICIT NONE
C
C-----
C      LOCAL VARIABLES
C-----
      INTEGER IALLCELL, IDD, NUMCELL, IL, I, K, IGRID
      INTEGER ICOUNT, ICM, ICP, IRM, IRP, ITRACK, IDD2, III, NTOTRCH, IC, IR
      INTEGER IDONE, IFLAG, IFLAG2, IFILL, IFILL2, OUTIT, OUTITMAX, INIT
      INTEGER INITMAX
      INTEGER II, J, HRUFLG, NUMACT, IDUM1, IDUM2
      REAL CLOSEZERO, PCT, TOL, DPIT !INCREMENT TO RAISE ELEVATION
C-----
C1-----ORDER ALL STORMFLOW CELLS BASED ON GRAVITY FLOW PATHS (TOP OF
      HILLS DOWN)
      !CASCADE_FILL SUBROUTINE
      !NOTE: ****CONSIDER MOVING THESE VARIABLES TO INPUT FILE***
      INITMAX=100000000
      IDONE=0
      DPIT=0.1
      TOL=0.5*DPIT
      OUTIT=0 !NUMBER OF OUTER ITERATIONS
      OUTITMAX=100000 !MAX NUMBER OF OUTER ITERATIONS
      INIT=0
      IF (IFILLFLG.EQ. 1) THEN
      DO WHILE (IDONE.EQ.0)
      !NOTE: IDONE=1 THE DEM COMPLETELY REPROCESSED AND NO SWALES EXIST
      IFLAG2=0
      !NOTE: NUMIT=0 ARRAY FOR STORING NUM OF ITERATIONS/CELL ***MOVE
      TO MODULE IF DECIDE TO KEEP***
      DO IR=1,NROW
      DO IC=1,NCOL
      IF (IBOUND(IC,IR,1).EQ.1) THEN
      IFLAG=1
      !NOTE: IFLAG=0 A LOWER CELL EXISTS AROUND CURRENT CELL
C
C2-----CHECK TO SEE IF CASCADE OPERATION WILL BE OUT OF BOUNDS
      IF (((IR+1).GT.NROW) .OR. ((IC+1).GT. NCOL)) THEN
      WRITE(IOUT,*)"CHECK MODEL DIMENSIONS-",
+      "EXITING CASCADE FILL"

```

```

EXIT
ELSE IF (((IR-1).LT.NROW) .OR. ((IC-1).LT. NCOL)) THEN
  WRITE(IOUT,*)"CHECK MODEL DIMENSIONS-",
  "EXITING CASCADE FILL"
EXIT
END IF
C
C3-----CASCADE CELL BY CELL FILL PROCEDURE
CHECKSWALE:DO WHILE(IFLAG.EQ.1)
  IF(STRMCELL(IC,IR).GT.0) EXIT CHECKSWALE
  !INSIDE CELLS
  IF ( ((IC.GT.1) .AND. (IC.LT.NCOL)) .AND. ((IR.GT.1)
  .AND. (IR.LT.NROW)) ) THEN
    IF ( (DEM2(IC,IR)-DEM2(IC,(IR-1)).GT.TOL).AND.
    (IBOUND(IC,(IR-1),1).NE. 0) ) THEN
      IFLAG=0
    ELSEIF ( (DEM2(IC,IR)-DEM2(IC,(IR+1)).GT.TOL).AND.
    (IBOUND(IC,(IR+1),1).NE. 0) ) THEN
      IFLAG=0
    ELSEIF ( (DEM2(IC,IR)-DEM2((IC-1),IR).GT.TOL).AND.
    (IBOUND((IC-1),IR,1).NE. 0) ) THEN
      IFLAG=0
    ELSE IF ( (DEM2(IC,IR)-DEM2((IC+1),IR).GT.TOL).AND.
    (IBOUND((IC+1),IR,1).NE. 0) ) THEN
      IFLAG=0
    END IF
  ENDIF
  !TOP LEFT CORNER
  IF ( ((IR.EQ.1) .AND. (IC.EQ.1)) ) THEN
    IF ( (DEM2(IC,IR)-DEM2(IC,(IR+1)).GT.TOL).AND.
    (IBOUND(IC,(IR+1),1).NE. 0) ) THEN
      IFLAG=0
    ELSE IF ( (DEM2(IC,IR)-DEM2((IC+1),IR).GT.TOL).AND.
    (IBOUND((IC+1),IR,1).NE. 0) ) THEN
      IFLAG=0
    END IF
  ENDIF
  !UPPER EDGE
  IF ( ((IC.NE.1) .AND. (IC.NE.NCOL)) .AND. (IR.EQ.1))
  THEN
    IF ( (DEM2(IC,IR)-DEM2(IC,(IR+1)).GT.TOL).AND.
    (IBOUND(IC,(IR+1),1).NE. 0) ) THEN
      IFLAG=0
    ELSEIF ( (DEM2(IC,IR)-DEM2((IC-1),IR).GT.TOL).AND.
    (IBOUND((IC-1),IR,1).NE. 0) ) THEN
      IFLAG=0
    ELSE IF ( (DEM2(IC,IR)-DEM2((IC+1),IR).GT.TOL).AND.
    (IBOUND((IC+1),IR,1).NE. 0) ) THEN
      IFLAG=0
    END IF
  ENDIF
  !TOP RIGHT CORNER
  IF ((IC.EQ.NCOL) .AND. (IR.EQ.1)) THEN
    IF ( (DEM2(IC,IR)-DEM2(IC,(IR+1)).GT.TOL).AND.
    (IBOUND(IC,(IR+1),1).NE. 0) ) THEN

```

```

        IFLAG=0
        ELSEIF ( (DEM2(IC,IR)-DEM2((IC-1),IR)).GT.TOL).AND.
+           (IBOUND((IC-1),IR,1).NE. 0) ) THEN
            IFLAG=0
        END IF
    ENDIF
    !LEFT HAND SIDE
    IF ((IC.EQ.1).AND.((IR.NE.1).AND.(IR.NE.NROW))) THEN
+       IF ( (DEM2(IC,IR)-DEM2(IC,(IR-1))).GT.TOL).AND.
+           (IBOUND(IC,(IR-1),1).NE. 0) ) THEN
            IFLAG=0
        ELSEIF ( (DEM2(IC,IR)-DEM2(IC,(IR+1))).GT.TOL).AND.
+           (IBOUND(IC,(IR+1),1).NE. 0) ) THEN
            IFLAG=0
        ELSE IF ( (DEM2(IC,IR)-DEM2((IC+1),IR)).GT.TOL).AND.
+           (IBOUND((IC+1),IR,1).NE. 0) ) THEN
            IFLAG=0
        END IF
    ENDIF
    !LOWER LEFT CORNER
    IF ((IC.EQ.1).AND.(IR.EQ.NROW)) THEN
+       IF ( (DEM2(IC,IR)-DEM2(IC,(IR-1))).GT.TOL).AND.
+           (IBOUND(IC,(IR-1),1).NE. 0) ) THEN
            IFLAG=0
        ELSE IF ( (DEM2(IC,IR)-DEM2((IC+1),IR)).GT.TOL).AND.
+           (IBOUND((IC+1),IR,1).NE. 0) ) THEN
            IFLAG=0
        END IF
    ENDIF
    !LOWER EDGE
    THEN
    IF (((IC.NE.1).AND.(IC.NE.NCOL)).AND.(IR.EQ.NROW))
+       IF ( (DEM2(IC,IR)-DEM2(IC,(IR-1))).GT.TOL).AND.
+           (IBOUND(IC,(IR-1),1).NE. 0) ) THEN
            IFLAG=0
        ELSEIF ( (DEM2(IC,IR)-DEM2((IC-1),IR)).GT.TOL).AND.
+           (IBOUND((IC-1),IR,1).NE. 0) ) THEN
            IFLAG=0
        ELSE IF ( (DEM2(IC,IR)-DEM2((IC+1),IR)).GT.TOL).AND.
+           (IBOUND((IC+1),IR,1).NE. 0) ) THEN
            IFLAG=0
        END IF
    ENDIF
    !LOWER RIGHT CORNER
    IF ((IC.EQ.NCOL).AND.(IR.EQ.NROW)) THEN
+       IF ( (DEM2(IC,IR)-DEM2(IC,(IR-1))).GT.TOL).AND.
+           (IBOUND(IC,(IR-1),1).NE. 0) ) THEN
            IFLAG=0
        ELSEIF ( (DEM2(IC,IR)-DEM2((IC-1),IR)).GT.TOL).AND.
+           (IBOUND((IC-1),IR,1).NE. 0) ) THEN
            IFLAG=0
        END IF
    ENDIF
    !RIGHT HAND SIDE

```

```

THEN
    IF ((IC.EQ.NCOL).AND.((IR.NE.1) .AND. (IR.NE.NROW)))
        IF ( (DEM2(IC,IR)-DEM2(IC,(IR-1)).GT.TOL).AND.
            (IBOUND(IC,(IR-1),1).NE. 0) ) THEN
            IFLAG=0
        ELSEIF ( (DEM2(IC,IR)-DEM2(IC,(IR+1)).GT.TOL).AND.
            (IBOUND(IC,(IR+1),1).NE. 0) ) THEN
            IFLAG=0
        ELSEIF ( (DEM2(IC,IR)-DEM2((IC-1),IR).GT.TOL).AND.
            (IBOUND((IC-1),IR,1).NE. 0) ) THEN
            IFLAG=0
        END IF
    ENDIF
    IF (IFLAG.EQ.1) THEN
        DEM2(IC,IR)=DEM2(IC,IR)+ DPIT
        IFLAG2=1 !NOTE: INDICATES THAT A CELL WAS UPDATED
        !NOTE: IDONE DOES NOT GET SET TO 1 UNTIL IFLAG2=0
        !NUMIT(IC,IR)=NUMIT(IC,IR)+1
        !NOTE: STORES NUMBER OF INCREMENTS FOR EACH CELL
    ENDIF
    INIT=INIT+1
    IF (INIT.GT.INITMAX) THEN
        PRINT*, "THERE WERE GREATER THAN", INITMAX,
            "INNER ITERATIONS...EXITING CASCADE_FILL"
        EXIT CHECKSWALE
    ENDIF
    END DO CHECKSWALE
    INIT=0
    ENDIF
    END DO
    OUTIT=OUTIT+1
    IF (OUTIT.GT.OUTITMAX) THEN
        PRINT*, "THERE WERE GREATER THAN", OUTITMAX,
            "OUTER ITERATIONS. RESETTING DEM TO ORIGINAL SURFACE.
            EXITING CASCADE_FILL."
        DO IR=1, NROW
            DO IC=1, NCOL
                DEM2(IC,IR)=BOTM(IC,IR,0)
            END DO
        END DO
        EXIT
    ENDIF
    IF (IFLAG2.EQ.0) THEN
        IDONE=1
    ENDIF
    END DO
ENDIF
C
C4-----WRITE OUT FILLED SURFACE TO OUTPUT FILE
WRITE (IOUT, *)"FILLED DEM FOR SURFACE CASCADES"
DO IR=1,NROW
    DO IC=1,NCOL
        WRITE (IOUT, '(50F7.2)') DEM2(IC,IR)
    END DO

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```

      END DO
C
C5-----BEGIN CASCADE EVALUATION
      IALLCELL=0
      IDD=1
      NUMCELL=NROW*NCOL
      CLOSEZERO = 1.0E-7
      I=1
      IL = 1
      DO IR = 1,NROW
        DO IC = 1,NCOL
          IF( IBOUND(IC, IR, IL).LE.0 ) THEN
            IORDERCHK(I)=-999
          END IF
          I=I+1
        END DO
      END DO
      II = 1
      IDD = 1
      ORDERCELL: DO WHILE( IALLCELL.EQ.0 )
        IF ( II.LE.NUMCELL ) THEN
          IDD2 = 1
          DO IR = 1,NROW
            DO IC = 1,NCOL
              ICOUNT=0
              IF(IORDERCHK(IDD2).EQ.0) THEN
                DO III = 3,10
                  IORDER(III,IDD) = 0
                END DO
              ICOUNT=0
              ICM=IC-1
              ICP=IC+1
              IRM=IR-1
              IRP=IR+1
C
C6-----DETERMINE THE NUMBER OF SURROUNDING CELLS THAT HAVE A
C-----LOWER ELEVATION.
              IF( IC.GT.1 ) THEN
                IF ( IBOUND(ICM, IR, 1).GT.0 ) THEN
                  IF(DEM2(IC, IR).GE.DEM2(ICM, IR)) THEN
                    ICOUNT=ICOUNT+1
                    IF(DEM2(IC, IR)-DEM2(ICM, IR).GT.CLOSEZERO)
                      IORDER(7,IDD)=1
                    ELSE IF(IORDERCHK(IDD2-1).GT.0) THEN
                      ICOUNT=ICOUNT+1
                    END IF
                    IF(DEM2(IC, IR).LT.DEM2(ICM, IR)) IORDER(3,IDD)=1
                  ELSE
                    ICOUNT=ICOUNT+1
                    IORDER(3,IDD)=0
                    IORDER(7,IDD)=0
                  END IF
                ELSE
                  ICOUNT=ICOUNT+1
                  IORDER(3,IDD)=0
                END IF
              END IF
            END DO
          END DO
        END WHILE
      END DO

```

```

IORDER(7,IDD)=0
END IF
IF( IC.LT.NCOL ) THEN
  IF ( IBOUND(ICP, IR, 1).GT.0 )THEN
    IF(DEM2(IC, IR).GE.DEM2(ICP, IR))THEN
      ICOUNT=ICOUNT+1
      IF(DEM2(IC, IR)-DEM2(ICP, IR).GT.CLOSEZERO)
        IORDER(8,IDD)=1
      ELSEIF( IORDERCHK( IDD2+1) .GT.0 )THEN
        ICOUNT=ICOUNT+1
      END IF
      IF(DEM2(IC, IR).LT.DEM2(ICP, IR))IORDER(4,IDD)=1
    ELSE
      ICOUNT=ICOUNT+1
      IORDER(4,IDD)=0
      IORDER(8,IDD)=0
    END IF
  ELSE
    ICOUNT=ICOUNT+1
    IORDER(4,IDD)=0
    IORDER(8,IDD)=0
  END IF
IF( IR.GT.1 ) THEN
  IF ( IBOUND(IC, IRM, 1).GT.0 )THEN
    IF(DEM2(IC, IR).GE.DEM2(IC, IRM))THEN
      ICOUNT=ICOUNT+1
      IF(DEM2(IC, IR)-DEM2(IC, IRM).GT.CLOSEZERO)
        IORDER(9,IDD)=1
      ELSEIF( IORDERCHK( IDD2-NCOL) .GT.0 )THEN
        ICOUNT=ICOUNT+1
      END IF
      IF(DEM2(IC, IR).LT.DEM2(IC, IRM))IORDER(5,IDD)=1
    ELSE
      ICOUNT=ICOUNT+1
      IORDER(5,IDD)=0
      IORDER(9,IDD)=0
    END IF
  ELSE
    ICOUNT=ICOUNT+1
    IORDER(5,IDD)=0
    IORDER(9,IDD)=0
  END IF
IF( IR.LT.NROW ) THEN
  IF ( IBOUND(IC, IRP, 1).GT.0 )THEN
    IF(DEM2(IC, IR).GE.DEM2(IC, IRP))THEN
      ICOUNT=ICOUNT+1
      IF(DEM2(IC, IR)-DEM2(IC, IRP).GT.CLOSEZERO)
        IORDER(10,IDD)=1
      ELSEIF( IORDERCHK( IDD2+NCOL) .GT.0 )THEN
        ICOUNT=ICOUNT+1
      END IF
      IF(DEM2(IC, IR).LT.DEM2(IC, IRP))IORDER(6,IDD)=1
    ELSE
      ICOUNT=ICOUNT+1
      IORDER(6,IDD)=0
    END IF

```

```

        IORDER(10,IDD)=0
        END IF
    ELSE
        ICOUNT=ICOUNT+1
        IORDER(6,IDD)=0
        IORDER(10,IDD)=0
    END IF
C
C7-----IF OUTFLOW TO ALL SURROUNDING CELLS HAS BEEN DETERMINED
C-----OR IF SURROUNDING CELLS HAVE LOWER ELEVATION.
        IF(ICOUNT.EQ.4)THEN
            IORDER(1,IDD)=IC
            IORDER(2,IDD)=IR
            IORDERCHK(IDD2) = 1
            IDD=IDD+1
        END IF
    END IF
    IDD2=IDD2+1
END DO
END DO
ITRACK=0
I = 1
FILTER: DO WHILE ( ITRACK.EQ.0 )
    IF ( I.LE.NUMCELL ) THEN
        IF(IORDER(1,I).EQ.0)THEN
            ITRACK=1
        END IF
    ELSE
        EXIT FILTER
    END IF
    I = I + 1
END DO FILTER
IF(ITRACK.EQ.0)IALLCCELL=1
ELSE
    EXIT ORDERCELL
END IF
II = II + 1
END DO ORDERCELL
NUMACTIVE = IDD - 1
END SUBROUTINE SOILZONE_ORDER
C
SUBROUTINE ACTIVECELL

USE GLOBAL, ONLY: IBOUND, NLAY, NCOL, NROW
USE GWFPSZMODULE, ONLY: IPSZBOUND
IMPLICIT NONE

-----
C LOCAL VARIABLES
-----
INTEGER::IC,IR, ILL
-----
C
C1-----NEW SUBROUTINE HERE TO FIND ACTIVE CELL AND ADD TO IPSZBOUND
ARRAY
    ILL = 1
    DO IR= 1, NROW

```

```

DO IC= 1, NCOL
  ILL=1
  ACTIVECHECK:DO WHILE ( ILL.LE.NLAY )
    IF ( IBOUND(IC, IR, ILL).NE.0 ) THEN
      EXIT
    END IF
    ILL = ILL + 1
  END DO ACTIVECHECK

  IF(ILL.LE.NLAY) IPSZBOUND(IC,IR) = ILL
END DO
END SUBROUTINE ACTIVECELL
C
SUBROUTINE ADDFINF(IUNITUZF,KKPER)
USE GLOBAL, ONLY: IBOUND, NLAY, NCOL, NROW, HNEW, BOTM, CV,RHS,
+ HCOF, DELR, DELC, IOUT
USE GWFPSZMODULE, ONLY: IPSZBOUND,R2SZ, UZF_REJ
USE GWFUZFMODULE, ONLY: FINF, IUZFBND, VKS, REJ_INF
IMPLICIT NONE
C LOCAL VARIABLES
C -----
INTEGER::IC, IR, IL, IUNITUZF, KKPER
DOUBLE PRECISION:: CELBOT, DELH
C -----
C1-----SUBROUTINE TO CALCULATE WATER BALANCE AND ADD TO FINF ARRAY IN
UZF
DO IR= 1, NROW
  DO IC= 1, NCOL
    IL=IPSZBOUND(IC, IR)
    IF (IL.GT.0) THEN
      IF (IL+1.LE.NLAY) THEN
        IF(IUZFBND(IC,IR).NE.0) THEN
!NOTE: THIS MAY PRESENT PROBLEMS WHEN CONSTANT HEAD IUZFBND SET TO ZERO
!WHICH MEANS ET, AND GROUNDWATER CONNECTION CANNOT BE ESTABLISHED IF
!GW HEAD AT BASE OF SZ FINF NOT ADDED TO UZF
          FINF(IC,IR)=R2SZ(IC,IR)/(DELC(IR)*DELR(IC))
          IF (FINF(IC,IR).GT.VKS(IC,IR)) THEN
            UZF_REJ(IC,IR)=(FINF(IC,IR)-VKS(IC,IR))
            FINF(IC,IR)=VKS(IC,IR)
            REJ_INF(IC,IR)=REJ_INF(IC,IR)+UZF_REJ(IC,IR)
          END IF
        END IF
      END IF
    END IF
  END DO
END DO
END SUBROUTINE ADDFINF
C
FUNCTION DDDHORIZ(HUP, TTOP, BBOT, IL)
USE GWFNWTMODULE
USE GWFUPWMODULE, ONLY: LAYTYPUPW
IMPLICIT NONE
C -----
-

```

```

C   SPECIFICATIONS:
C   -----
C   ARGUMENTS
C   -----
C   DOUBLE PRECISION HUP, TTOP, BBOT
C   DOUBLE PRECISION DDDHORIZ
C   -----
C   LOCAL VARIABLES
C   -----
C   DOUBLE PRECISION FACTOR, X, S, V, COF1, COF2, EPS, ACOF, Y
C   DOUBLE PRECISION EPSQD, Z
C   INTEGER IL
C   -----
C1----RETURNS DERIVATIVE OF HORIZONTAL CONDUCTANCE BASED ON SMOOTH
      FUNCTION
C   FUNCTION IS CALCULATED IN UPW PACKAGE IN SUBROUTINE SAT_THICK
      DDDHORIZ = 0.0D0
      IF ( LAYTYPUPW(IL).LE.0 ) RETURN
C2----STRAIGHT LINE WITH PARABOLIC SMOOTHING
      EPS = THICKFACT
      ACOF = 1.0 / (1.0 - EPS)
      X = (HUP-BBOT)/(TTOP-BBOT)
      IF ( X.LT.1.0D-9 ) X = 1.0D-9
      IF ( X.LT.0.0 ) THEN
          Y = 0.0
      ELSEIF(X.LT.EPS)THEN
          Y = ACOF * X / (EPS*(TTOP-BBOT))
      ELSEIF(X.LT.1.0D0-EPS)THEN
          Y = ACOF / (TTOP-BBOT)
      ELSEIF(X.LT.1.0D0)THEN
          X = 1.0 - X
          Y = - ACOF * X / (EPS * (TTOP - BBOT))
          Y = 1.0-Y
      ELSE
          Y = 0.0
      ENDIF
      FACTOR = Y
      DDDHORIZ = FACTOR
      END FUNCTION DDDHORIZ
C   -----

```

APPENDIX B: SZF CODE IMPLEMENTED INTO STREAM FLOW ROUTING PACKAGE.

```

      DOUBLE PRECISION FUNCTION SEEPFACE(L,DEP, IC, IR)
      USE GLOBAL,          ONLY: NCOL, NROW, NLAY, IOUT, DELR, DELC, CR,
+      IBOUND, BOTM, HCOF, RHS, CC, CV, HNEW
      USE GWFPSZMODULE,   ONLY: IPSZBOUND, STRMCELL, STRMCHD,
+      QSTORM
      USE GWFSFRMODULE,   ONLY: NSTRM, ISTRM, STRM, HSTRM
      USE GWFNWTMODULE,   ONLY: ICELL, IA, A
      USE GWFUPWMODULE,   ONLY: HKUPW, SN

!-----
      IMPLICIT NONE
      EXTERNAL DDHORIZ

C-----
      ! ARGUMENTS
C-----
      DOUBLE PRECISION HH, TTOP, BBOT, THICK, CONDIF, DH,
+      DDHORIZ, SSTRLEN, CHECK, SSTRTOP, STAGELEV, DEP, DUM,
+      DUM2
      REAL DUM3
      INTEGER IC, IR, IL, L, IJ, N

C-----
C1-----THIS SUBROUTINE SETS EACH STREAM CELL TO HAVE HEAD DEPENDENT
      BOUNDARIES TO ALLOW STAGE DEPENDENT FLOW FROM SOIL ZONE INTO
      STREAM
      DUM3=0.0
      CHECK=0.0
      IL=IPSZBOUND(IC,IR)
      STRMCELL(IC,IR)=L !NOTE: NEED TO MOVE STMT TO SZF
      SSTRTOP = STRM(3, L)
      STAGELEV=DEP+SSTRTOP
      DUM=HNEW(IC,IR,2)
      DO IL=1,1
      IF (STAGELEV.LT.HNEW(IC,IR,IL)) THEN
      IF (IBOUND(IC, IR, IL).GT.0)THEN
      IF (HNEW(IC,IR,IL).GT.BOTM(IC,IR,IL)) THEN
      IF (STAGELEV.GT.BOTM(IC,IR,IL)) THEN
      HH= HNEW(IC,IR,IL)
      TTOP=(BOTM(IC,IR,IL-1))
      BBOT=STAGELEV
      THICK=TTOP-BBOT
      SSTRLEN=2*STRM(1,L)
      CONDIF=(HKUPW(IC,IR,IL)*SSTRLEN)/((DELC(IR)+DELR(IC))/4)
      DH=DDHORIZ(HH,TTOP,BBOT,IL) !DSN/DH
      IJ=ICELL(IC,IR,IL)
      CHECK=(CONDIF*SN(IJ)*THICK*(HH-STAGELEV))
      ELSE
      HH= HNEW(IC,IR,IL)
      TTOP=(BOTM(IC,IR,IL-1))
      BBOT=(BOTM(IC,IR,IL))
      THICK=TTOP-BBOT
      SSTRLEN=2*STRM(1,L)
      CONDIF= (HKUPW(IC,IR,IL)*SSTRLEN)/((DELC(IR)+DELR(IC))/4)

```

```

        DH=DDHORIZ(HH,TTOP,BBOT,IL) !DSN/DH
        IJ=ICELL(IC,IR,IL)
        CHECK=(CONDIF*SN(IJ)*THICK*(HH-BBOT))
        END IF
    END IF
    END IF
    END IF
    DUM3=DUM3+CHECK
    END DO
    SEEPFACE=DUM3
END FUNCTION SEEPFACE
C-----
SUBROUTINE SEEPFACENWT(L, DEPNWT,KKPER, KKSTP)
C
    USE GLOBAL,          ONLY: NCOL, NROW, NLAY, IOUT, DELR, DELC, CR,
+   IBOUND, BOTM, HCOF, RHS, CC, CV, HNEW
    USE GWFPSZMODULE,   ONLY: IPSZBOUND, STRMCELL, STRMCHD,
+   QSTORM, HITR
    USE GWFSFRMODULE,   ONLY: NSTRM, ISTRM, STRM, HSTRM
    USE GWFNWTMODULE,   ONLY: ICELL, IA, A, BB
    USE GWFUPWMODULE,   ONLY: HKUPW, SN
C-----
    IMPLICIT NONE
    EXTERNAL DDHORIZ
C-----
    !ARGUMENTS
C-----
    DOUBLE PRECISION DELH, HH, TTOP, BBOT, THICK, CONDIF, DH, DCDH,
+   DDHORIZ, SSTRLEN, CHECK2, SSSTRTOP,STAGELEV, FINTERFLOW, DEPNWT,
+   DUMM1,DUMM2, FIFCHECK, DUMM3, DUMM4,DUMM5, HOLD,DBDH, AO, RHSO
    REAL DUMM6
    INTEGER IC, IR,IL, L, IJ, KKPER, KKSTP
C-----
C1-----THIS SUBROUTINE UPDATES THE RHS AND A NEWTON MATRIX WITH SZ
INTERFLOW VALUES
    CHECK2=0.0
    DUMM6=0.0
    FIFCHECK=0.0
    SSSTRTOP = STRM(3, L)
    STAGELEV=DEPNWT+SSSTRTOP
    IC=ISTRM(3, L)
    IR=ISTRM(2, L)
    DO IL=1,1 !NOTE: IF.GT.1 WILL ALLOW FOR LAYERS IN SZ
        IF (STAGELEV.LT.HNEW(IC,IR,IL)) THEN
            IF (IBOUND(IC, IR, IL).NE.0)THEN
                IF (HNEW(IC,IR,IL).GT.BOTM(IC,IR,IL))THEN
                    IF (STAGELEV.GT.BOTM(IC,IR,IL)) THEN
                        BBOT=STAGELEV
                    ELSE
                        BBOT=(BOTM(IC,IR,IL))
                    END IF
                ELSE
                    HH= HNEW(IC,IR,IL)
                    TTOP=(BOTM(IC,IR,IL-1))
                    BBOT=(BOTM(IC,IR,IL))
                    DBDH=0.0
                END IF
            END IF
        END IF
    END DO

```

```

THICK=TTOP-BBOT
SSTRLEN=2*STRM(1,L)
CONDIF= (HKUPW(IC,IR,IL)*SSTRLEN)/
+      ((DELC(IR)+DELR(IC))/4.0)
DH=DDHORIZ(HH,TTOP,BBOT,IL) !DSN/DH
IJ=ICELL(IC,IR,IL)
CHECK2=(CONDIF*SN(IJ)*THICK*(BBOT-HH))
      DUMM1=A(IA(IJ))
      DUMM2=RHS(IC,IR,IL)
DUMM3=(CONDIF*DH*THICK*(BBOT-HH))-(CONDIF*SN(IJ)*THICK)
+ DUMM4=(CONDIF*DH*THICK*BBOT)-(CONDIF*SN(IJ)*THICK)-
      (CONDIF*DH*THICK*DUMM5)
      !NOTE: DUMM4=DUMM3 PREV ITER; DUMM5=HOLD
      A(IA(IJ))= A(IA(IJ))+DUMM3
      BB(IJ)= BB(IJ) + DUMM3*HH
      RHS(IC,IR,IL)= RHS(IC,IR,IL) - CHECK2
      DUMM5=DUMM3*(HH-HOLD)+CHECK2
      !NOTE: DUMM5 SHOULD APPROACH FINTERFLOW
FIFCHECK= (A(IA(IJ))-DUMM1)*HNEW(IC,IR,IL)-
      RHS(IC,IR,IL)
DUMM6=DUMM6+CHECK2
      END IF
      END IF
      END IF
      END DO
QSTORM(IC,IR)= QSTORM(IC,IR)+DUMM6
      IF (QSTORM(IC,IR).GT.0.0) QSTORM(IC,IR)= 0.0
      !NOTE: THIS PREVENTS FLOW BACK INTO SZ
      HITR=HNEW
      END SUBROUTINE SEEPFACENWT

```